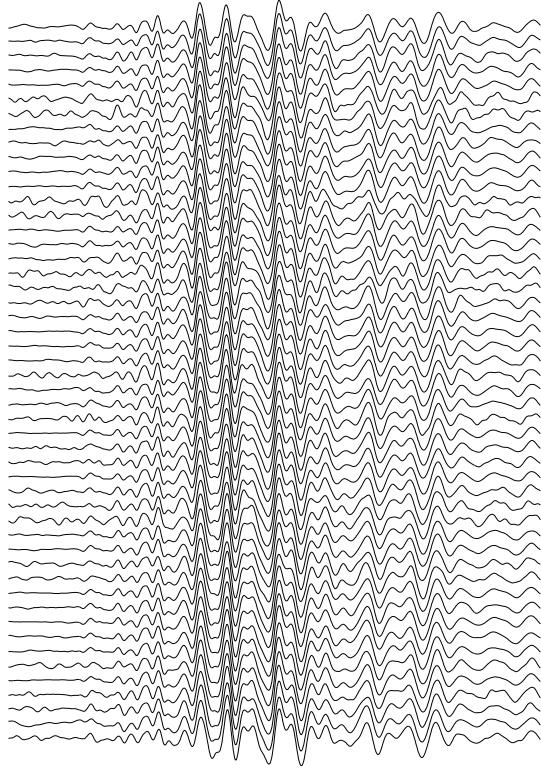


# Repeating icequakes and landquakes



Agnès Helmstetter

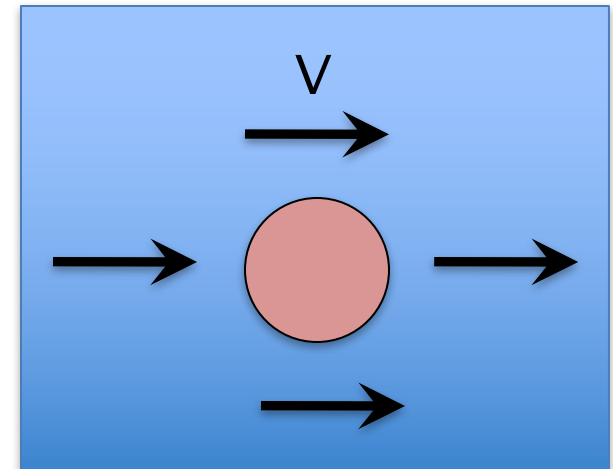
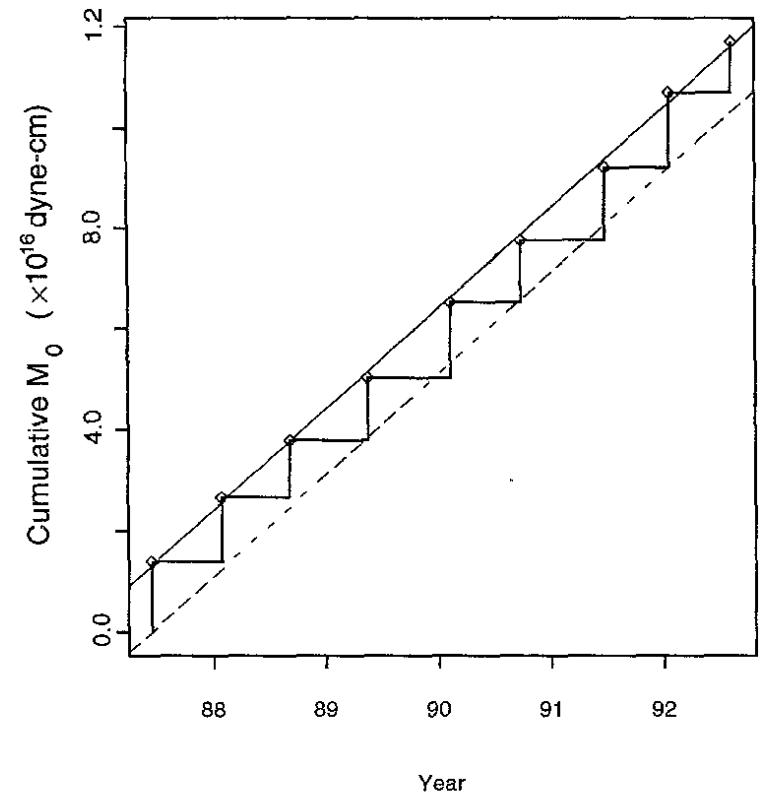
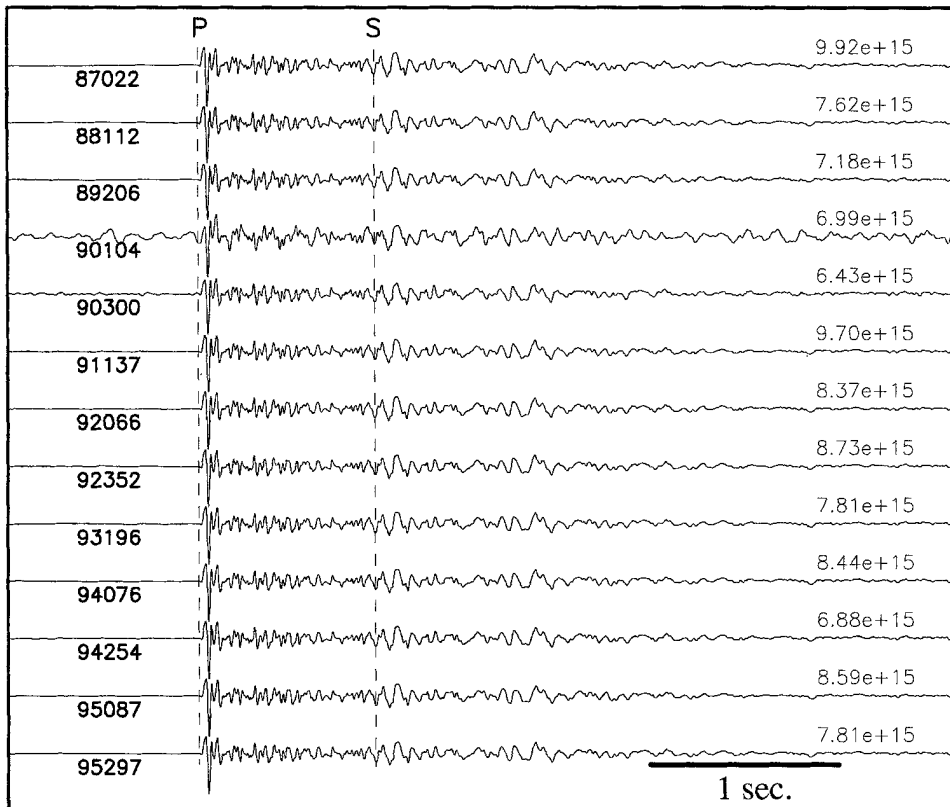
ISTerre, Université Grenoble-Alpes

Collaboration with Eric Larose, Laurent Baillet, Raphaël Mayoraz, Claudia Rössli, Fabian Walter, Barbara Nicolas, Michel Gay, Pierre Common



# Repeating earthquakes

Parkfield, California  
(Nadeau and Johnson, 1998)

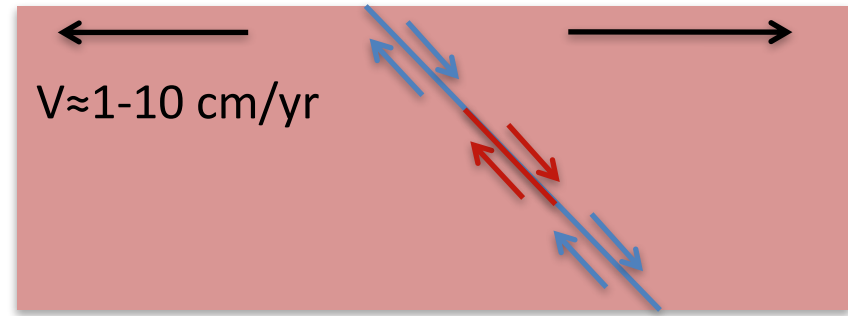


# Repeating earth/ice/land-quakes

## Repeating earthquakes

Tectonic loading

Asperities on the fault  
surrounded by creeping zones  
(# friction parameters?)

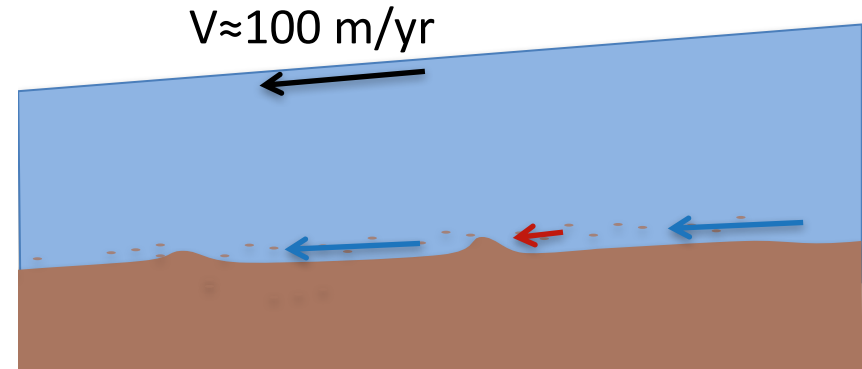


## Repeating icequakes

Slip at the ice/bed interface

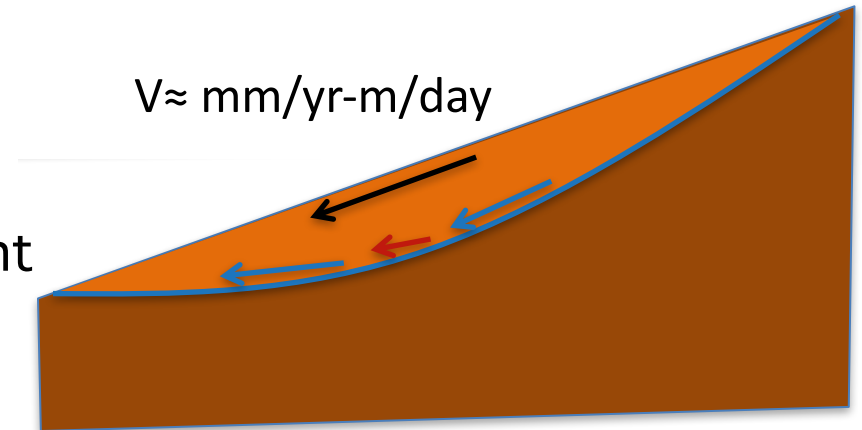
Loading by glacier movement

“sticky spots” ; rock debris? Bumps?



## Repeating landquakes

Loading = gravity + rockslide movement  
asperities : interface rugosity?



# Motivations

- > Proxy for aseismic slip?
- > Reproducible source
  - => can be used to detect small changes in the medium?
- > Precursor of a large rupture?



# Why landslides and glaciers?

Faster displacement, (usually) smaller size

⇒ longer time series

Shallower

⇒ easier to detect seismic events and to measure aseismic slip

Time dependent loading and forcings

⇒ analyze the influence of loading and forcings

⇒ more complex!

- daily and seasonal variations of velocity by a factor  $> 2$
- increase in water pressure
- ice : near melting point ⇒ change in rheology

# Outline

- > Icequakes

  - Antarctica, Greenland, Alpine glacier

- > Landslides

  - Rausu (Japan) and Aletsch (Switzerland)

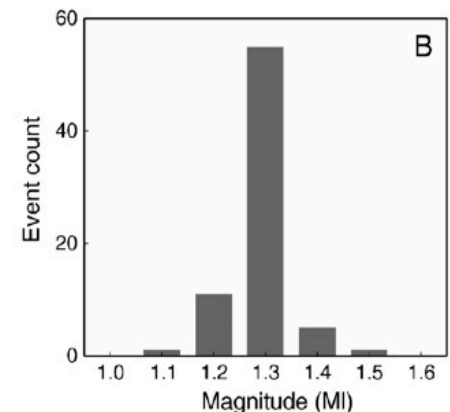
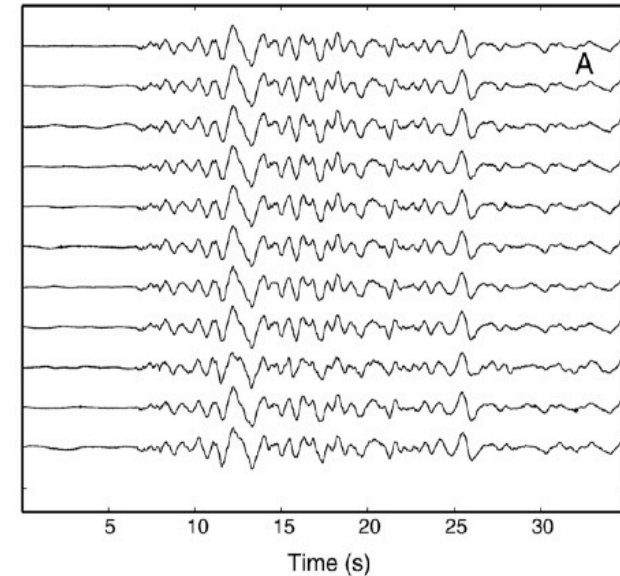
- > Rock-glacier

  - Gugla (Switzerland) and comparison with Mount Rainier (US)

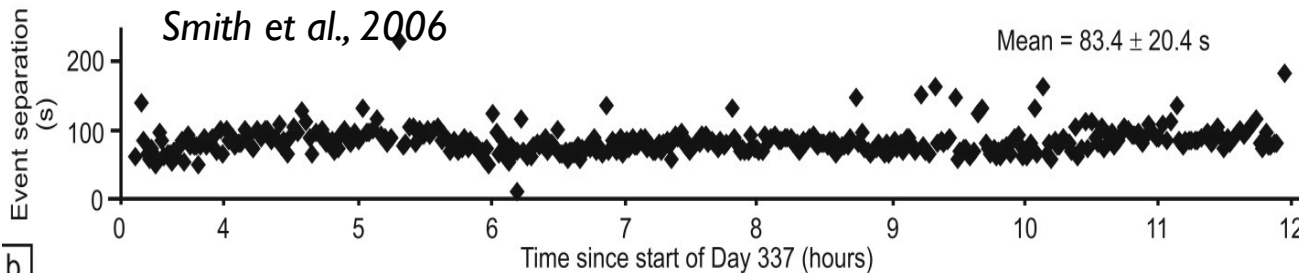
# Repeating icequakes in Antarctica

- > first observed by Anandakrishnan (1993)
- > different scales in magnitude, space and time
  - $M = -1.8$ ,  $dt = 83$  s (*Smith et al 2006*)
  - $M = 1.8$ ,  $dt = 25$  mn (*Zoet et al, 2012*)
  - $M_w = 7$ ,  $L = 200$  km, slip = 50 cm,  $dt = 12$  hr (*Wiens et al 2008*)

*Danesi, 2007*



*Smith et al., 2006*

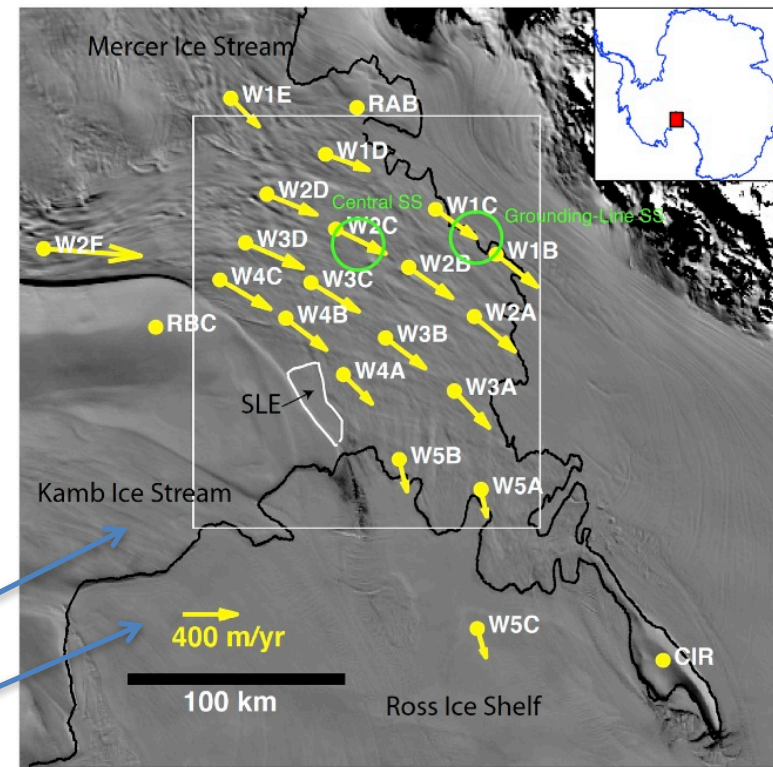


# Whillans Ice Stream

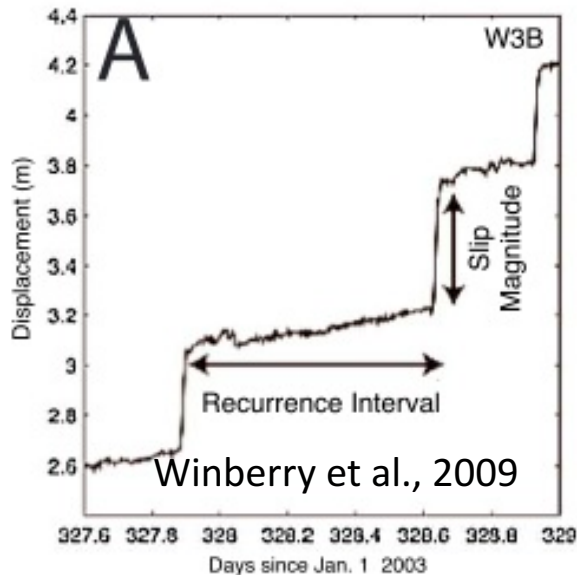
- > 2 events/day at high and low tides
- > rupture length  $\approx 200$  km
- > slip  $\approx 50$  cm,  $M_w=7$
- > duration 30 mn
- > rupture propagation 150 m/s
- > slip velocity  $\approx$  mm/s

Grounded ice

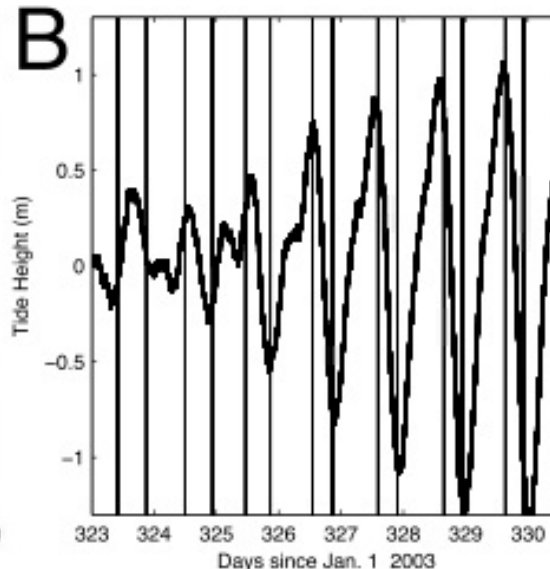
Floating ice



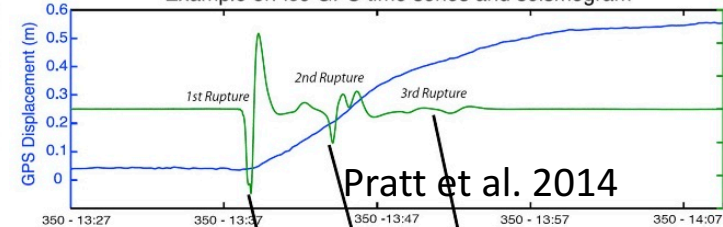
Displacement (GPS)



Tidal height and icequakes



Example on-ice GPS time series and seismogram



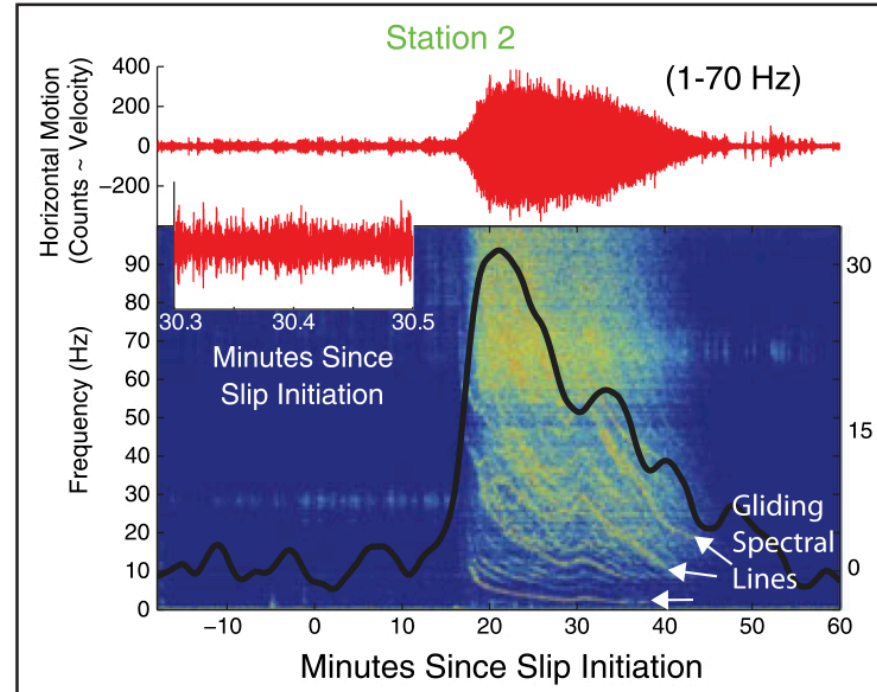
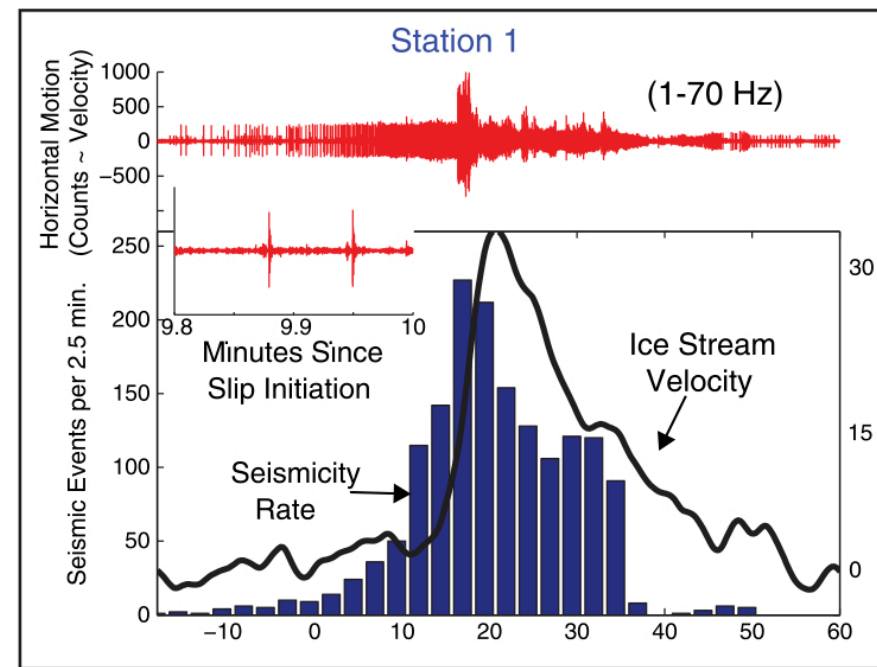
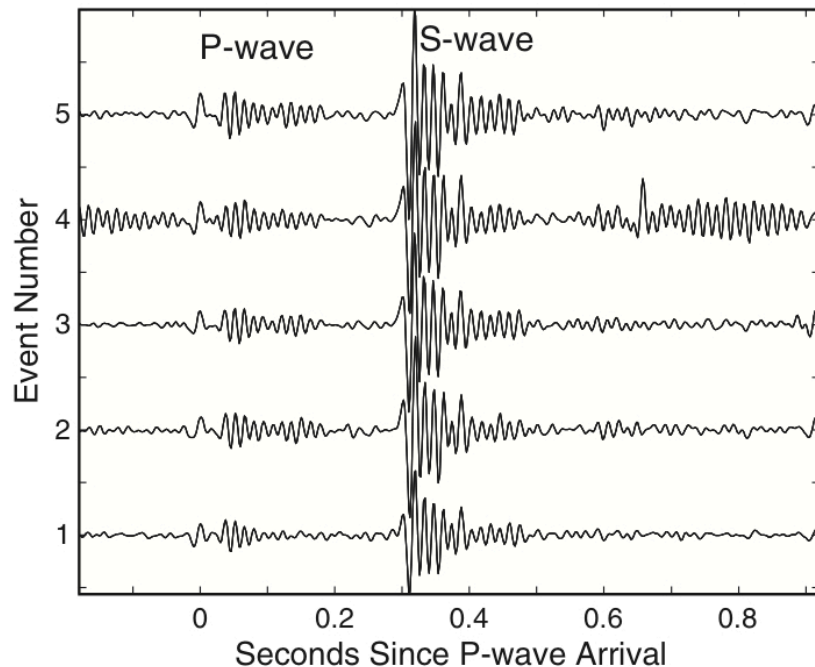
# Small repeaters occurring during rupture!

(Winberry et al, GRL 2013)

> 2 # behaviors at # stations

- repeating small icequakes  $\Delta t=5s$ , with rate slip rate
- or gliding tremor, with  $f_0 = 1 / \text{recurrence time}$

Micrearthquake Family Recorded at Station 1





# Repeating icequakes, Glacier d'Argentière

Mont-Blanc massif

Temperate glacier ( $T=0^{\circ}\text{C}$ )

3400-1600 m elevation

10 km in length

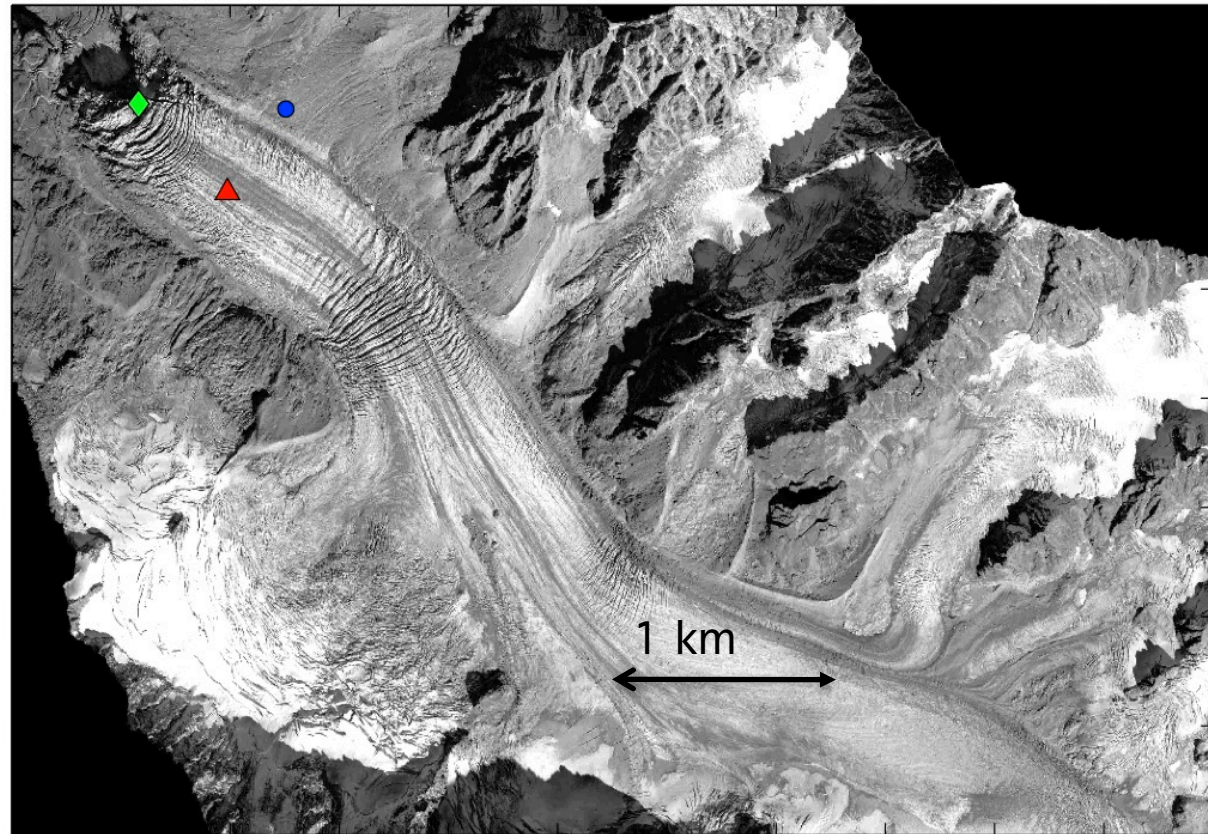
Velocity  $\approx 80$  m/yr

Thickness  $\approx 200$  m

(*Helmstetter et al, 2015*)

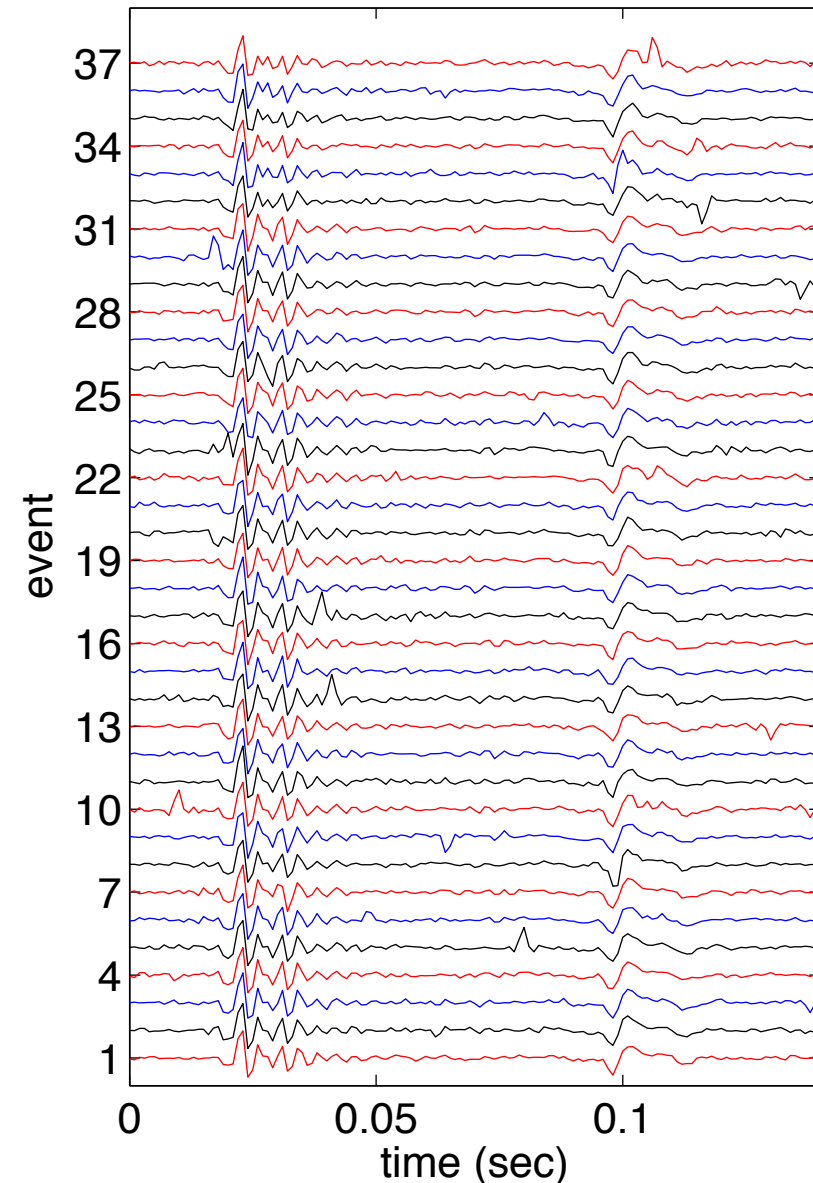
One 3C accelerometer

Sampling rate 1000 Hz



# Repeating icequakes, Glacier d'Argentière

- > similar waveforms
- > high-frequency  $\approx 200$  Hz
- > distinct P and S waves  $\Rightarrow$  deep source
- > 18 # clusters with 10-884 events
- > identification of template waveforms by STA/LTA
- > identification of similar events by template matching



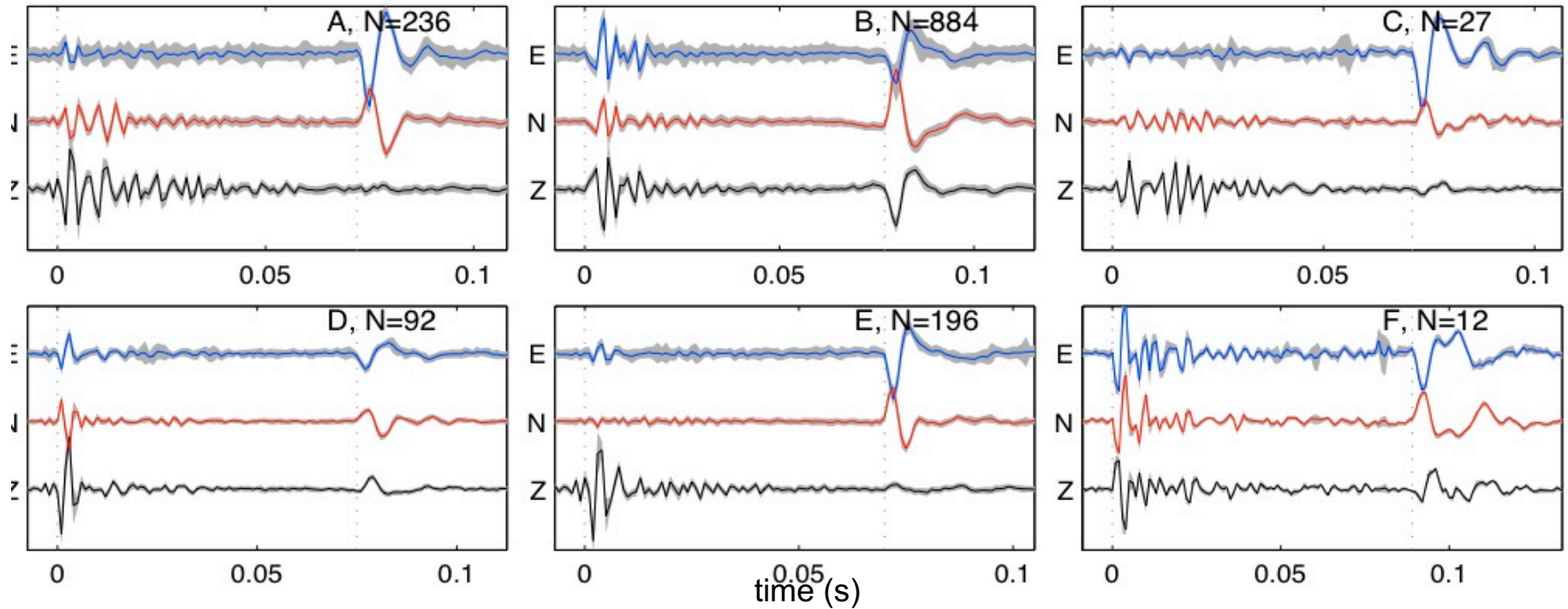


# Repeating icequakes, Glacier d'Argentière

Seismograms 3C (acceleration)

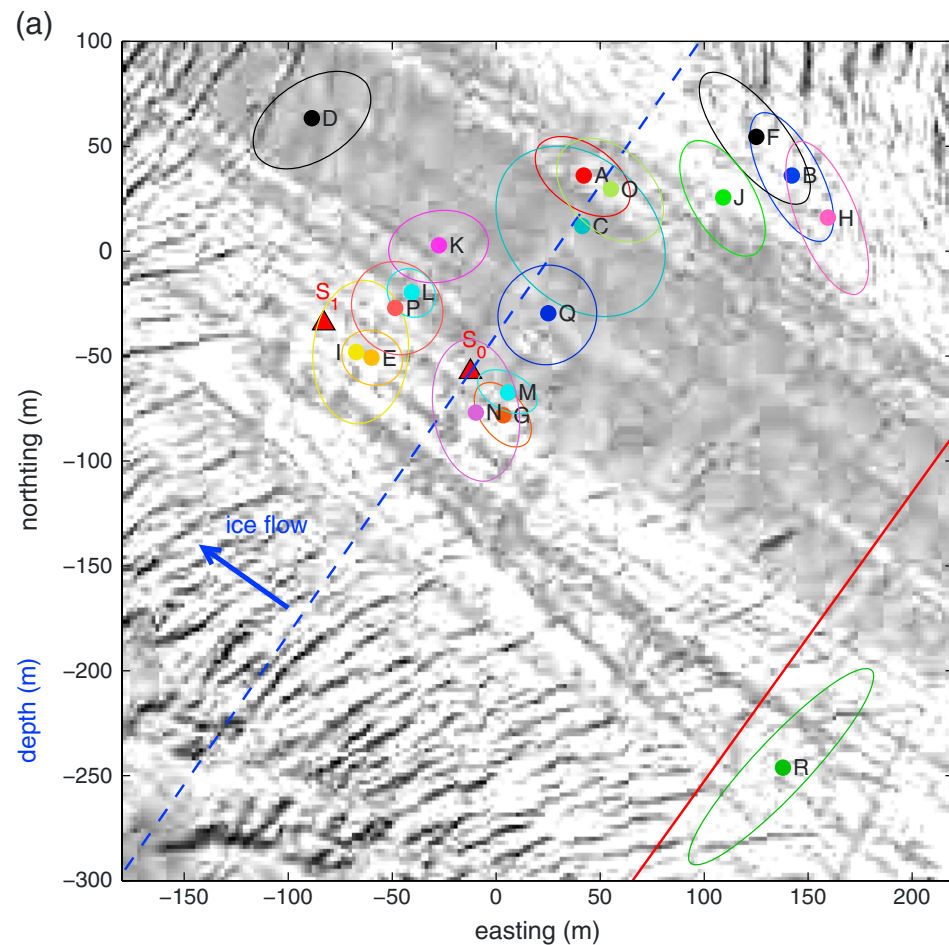
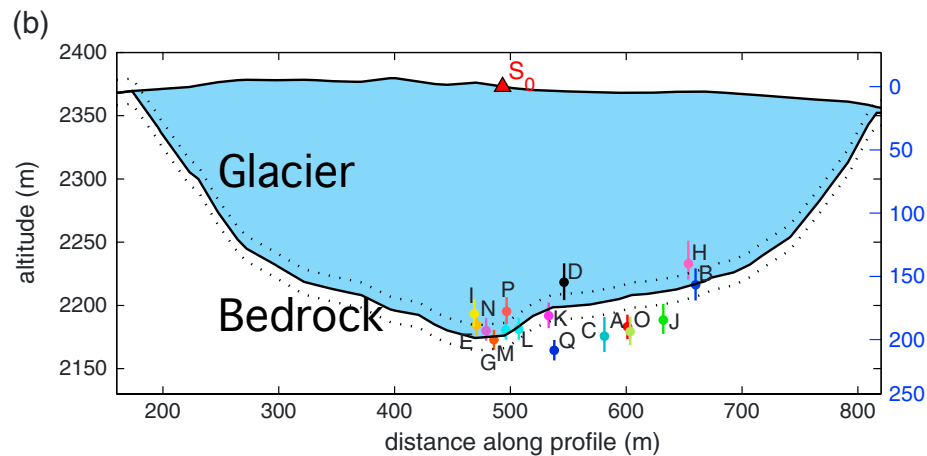
Stack of signals after normalization and alignment

Grey line : standard deviation among all events



# Location of repeating icequakes

- > Location with a single 3C sensor from P wave polarization and P-S time delay
- > Interface glacier – bedrock

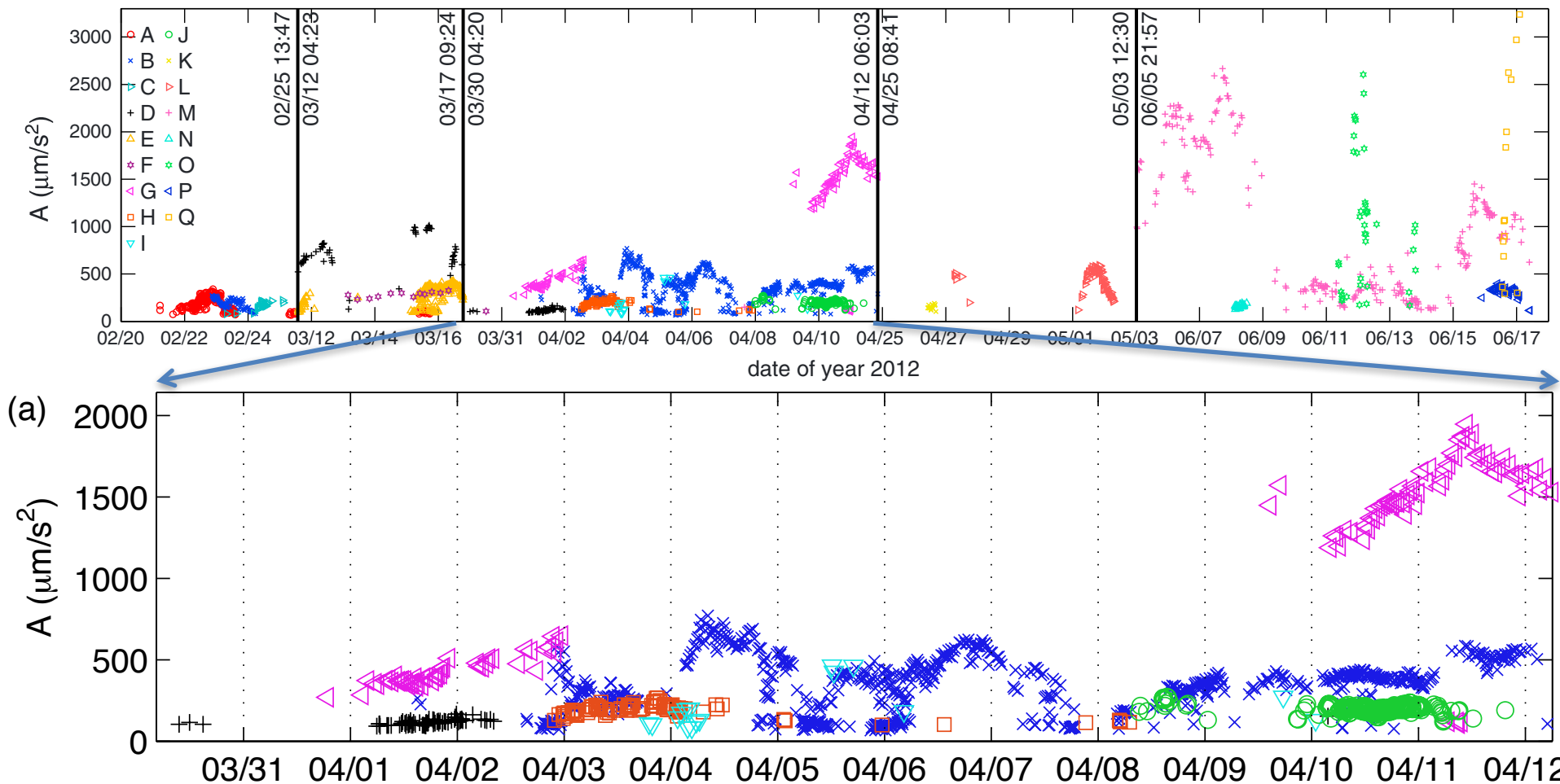


# Time series of icequakes amplitudes

February – June 2012 with large gaps (vertical lines)

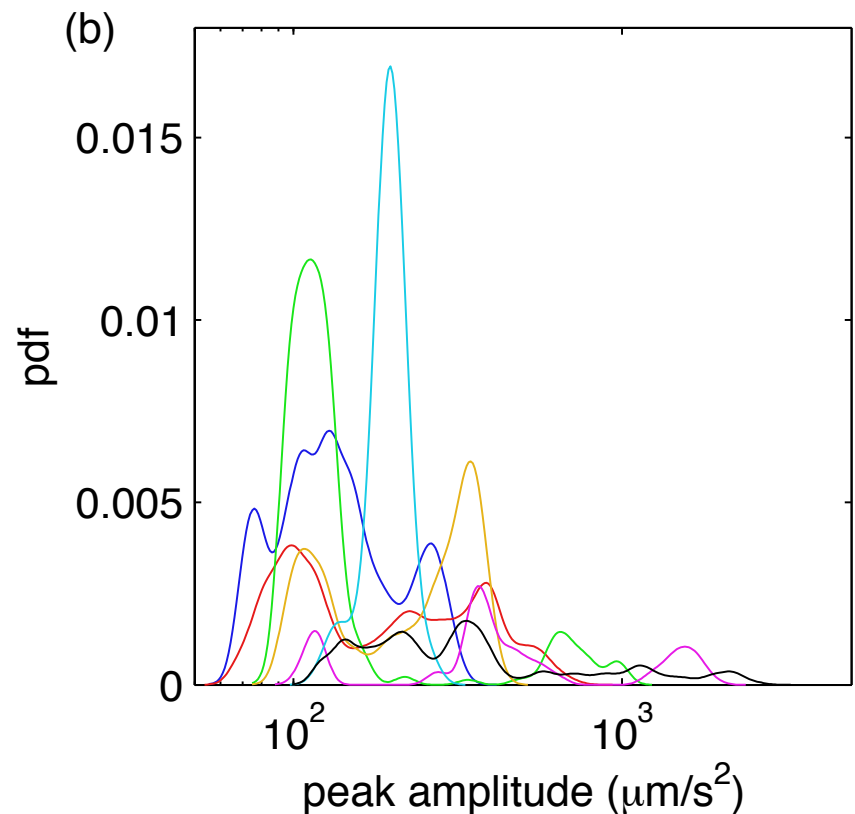
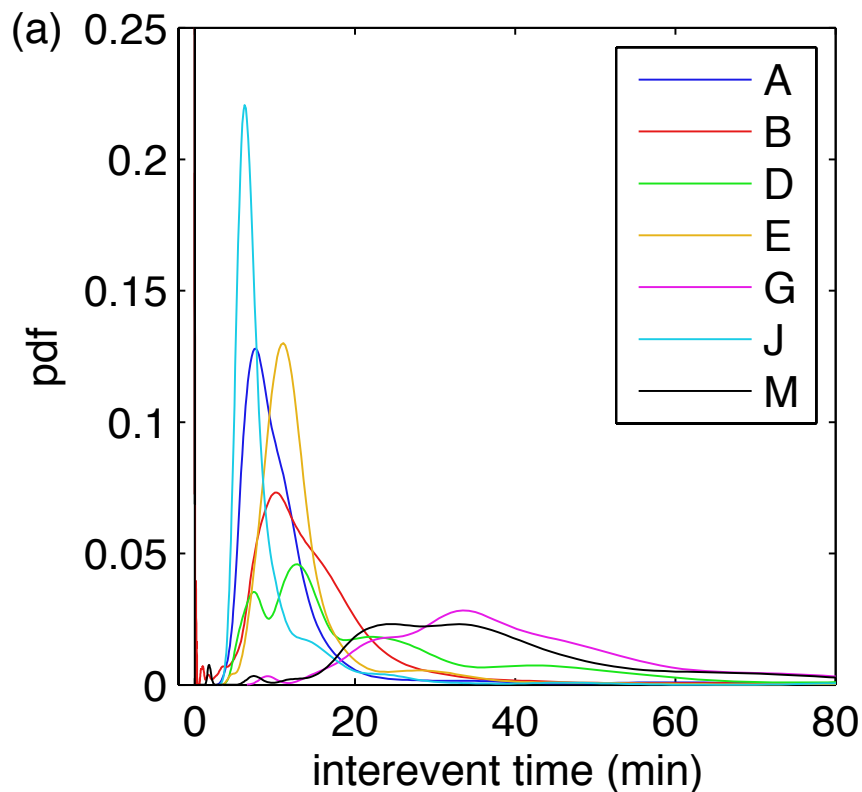
Succession of active and quiet periods

Progressive evolution of amplitude and recurrence times



# Distribution of inter-ev. times and amplitudes

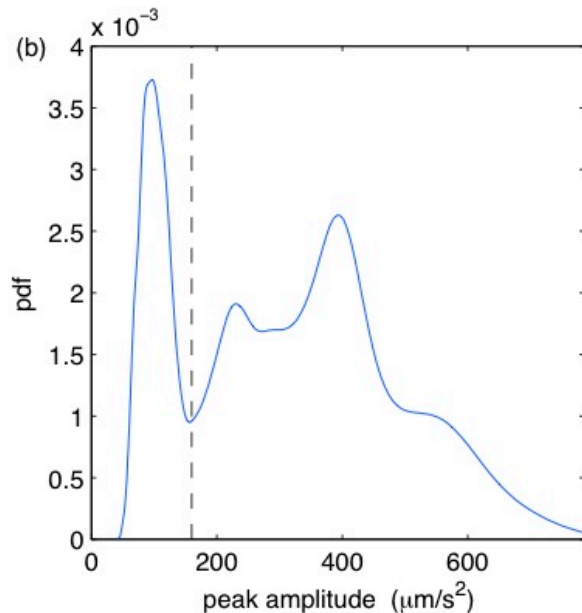
- > Distribution of inter-ev.times: quasi periodic  $\langle dt \rangle \approx 10$  mn
- > Distribution of amplitudes: much narrower than GR law



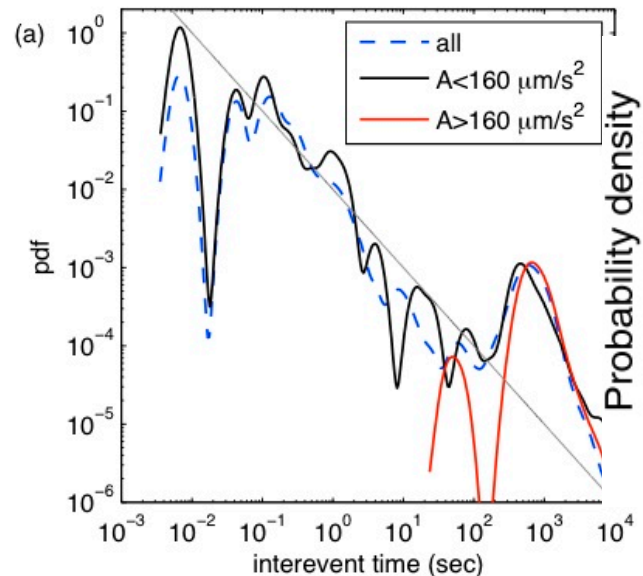
# Distribution of inter-ev. times and amplitudes

- > distribution of inter-ev.times: quasi periodic  $\langle dt \rangle \approx 10$  mn
- > and clustering for short dt and small A with power-law pdf(dt)
- > ... same as repeating earthquakes in Parkfield (*Lengliné et al; 2008*)

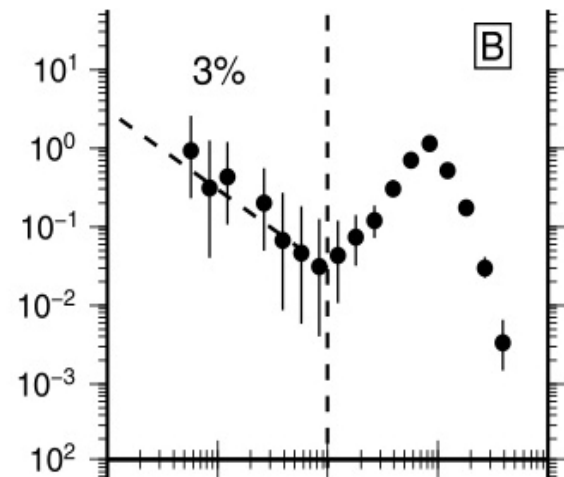
pdf(amplitude)  
Argentière



pdf(inter-ev.times)  
Argentière



pdf(inter-ev.time)  
Parkfield

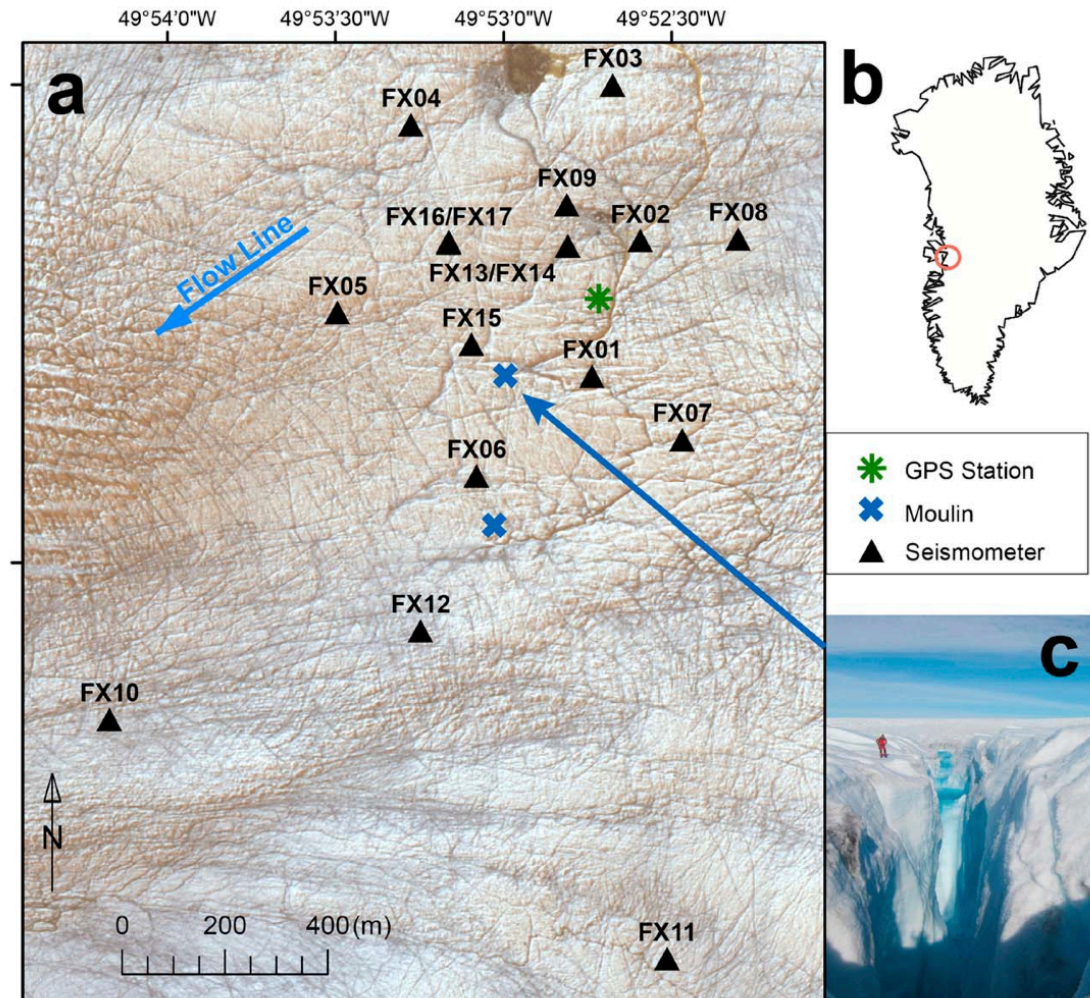


(*Lengliné et al, 2008*)



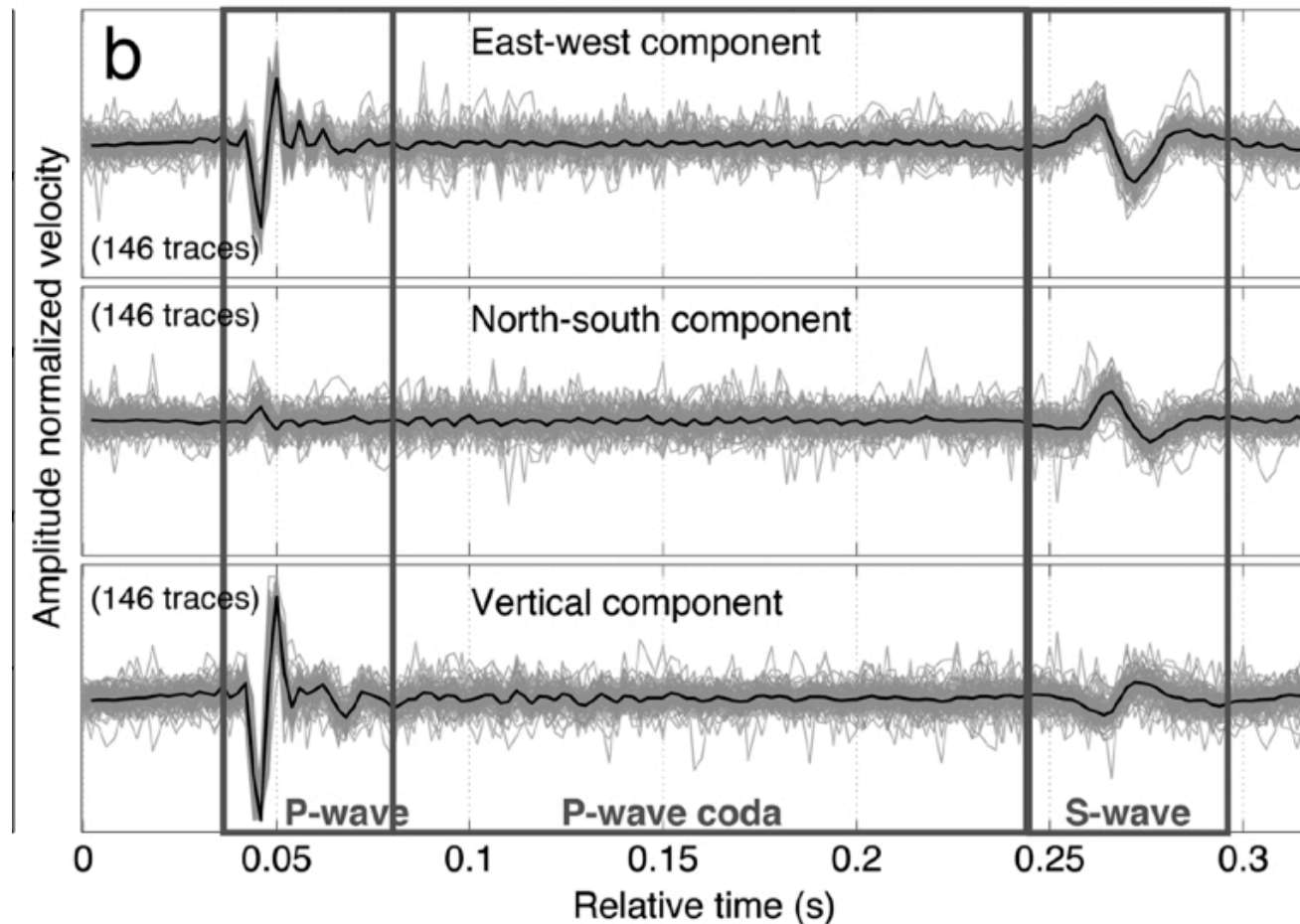
# Greenland

- > 3 July -15 August
- > 17 3C seismometers at the surface and in boreholes
- > Measurements of displacement, water level and water pressure in a moulin
- > Ice thickness  $\approx 650$  m
- > velocity  $\approx 27$  cm/day  
(Röösli et al, JGR 2016)



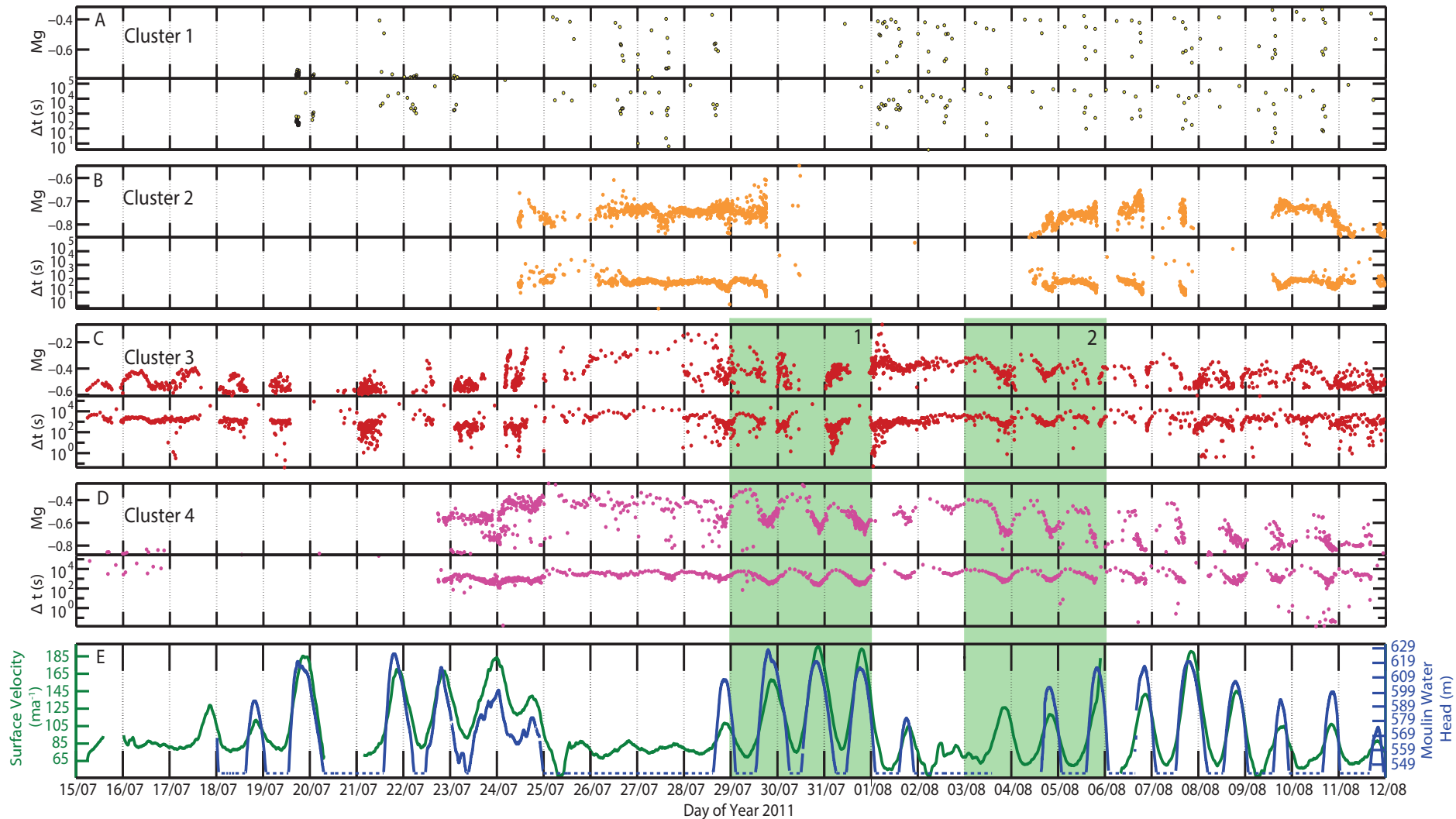
# Repeating icequakes in Greenland

More than 11,000 basal icequakes detected, grouped into > 100 clusters  
(*Roosli et al, JGR 2016*)

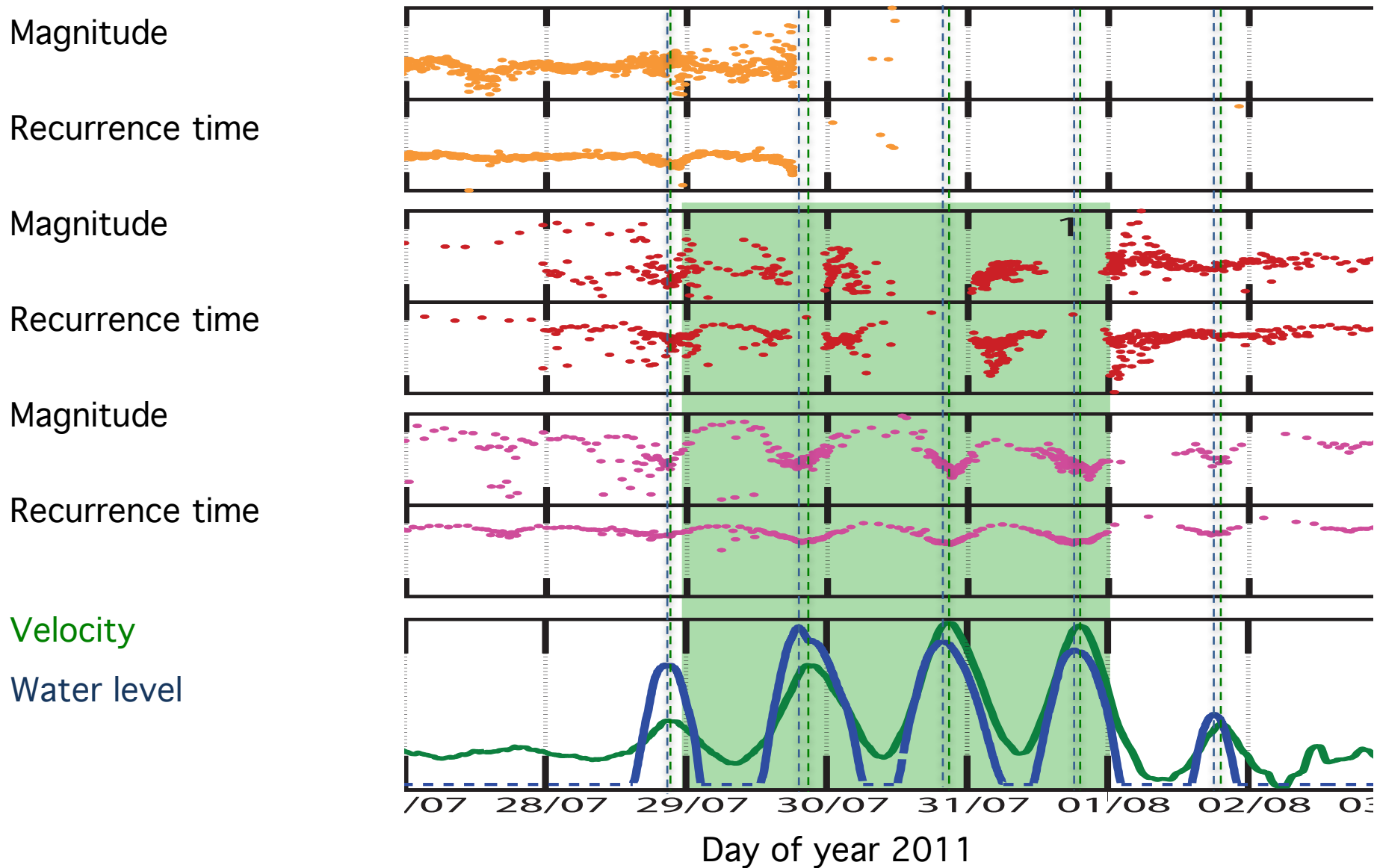




# Time series of icequakes magnitude, inter-event times, displacement and water level



# Time series ... zoom



# Time series : seismic and aseismic slip

- seismic slip  $\sim$  signal amplitude, cumulated over time bins of 0.1 day



## RESEARCH LETTER

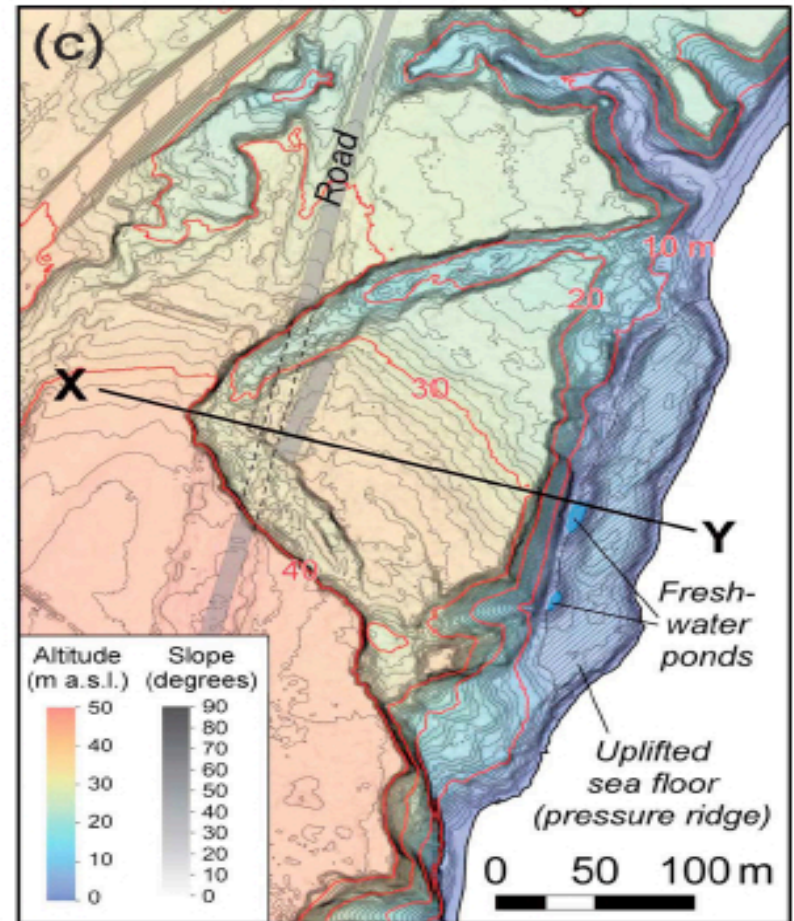
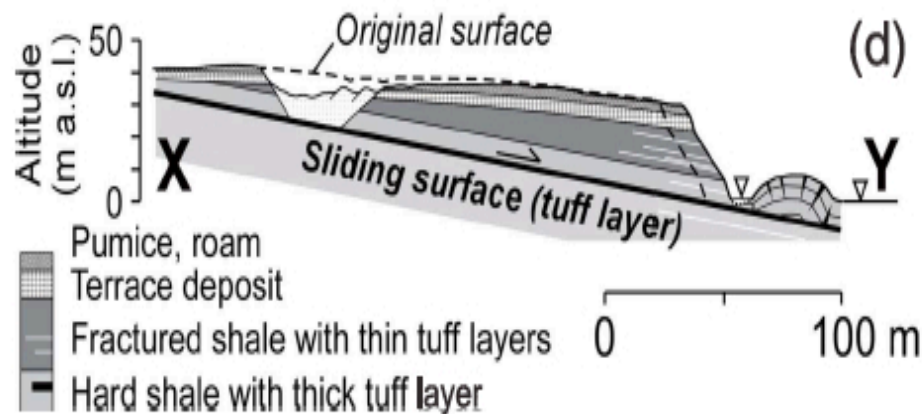
10.1002/2016GL069288

## Possible stick-slip behavior before the Rausu landslide inferred from repeating seismic events

### Key Points:

- We found tiny repeating earthquakes

Masumi Yamada<sup>1</sup>, Jim Mori<sup>1</sup>, and Yuki Matsushi<sup>1</sup>



## RESEARCH LETTER

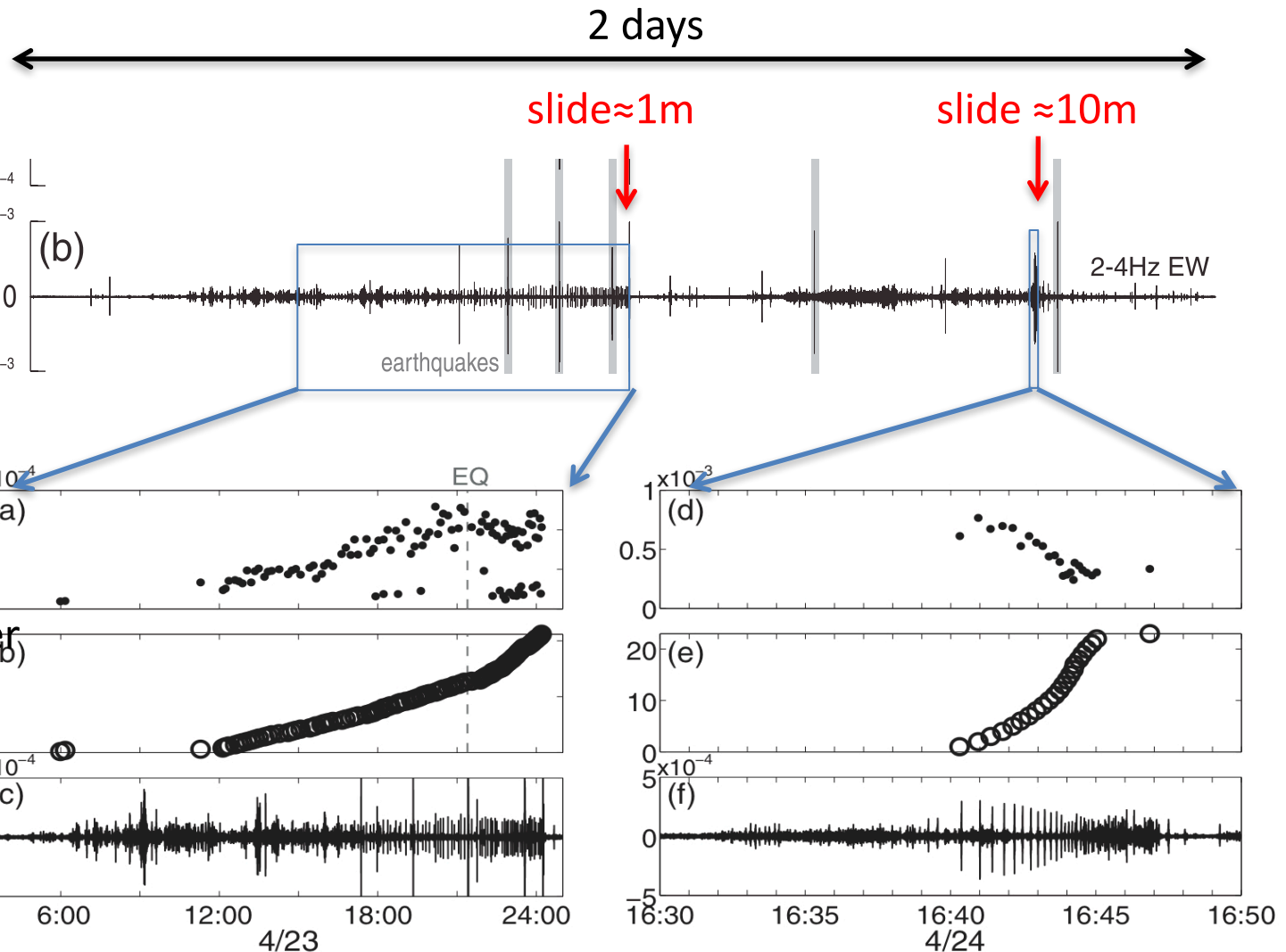
10.1002/2016GL069288

## Possible stick-slip behavior before the Rausu landslide inferred from repeating seismic events

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Masumi Yamada<sup>1</sup>, Jim Mori<sup>1</sup>, and Yuki Matsushi<sup>1</sup>





# Aletsch

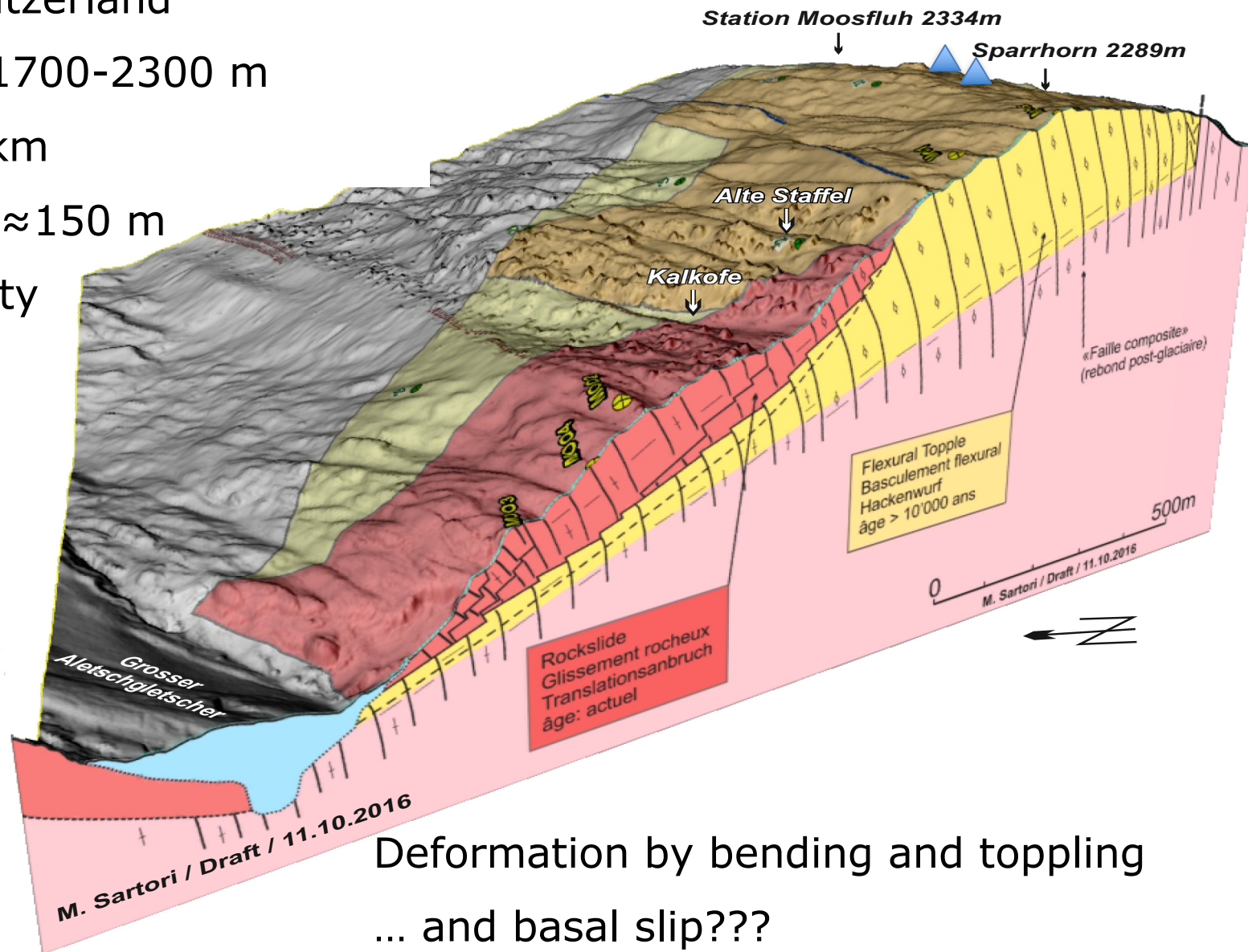
Valais, Switzerland

Elevation 1700-2300 m

Length  $\approx 1$  km

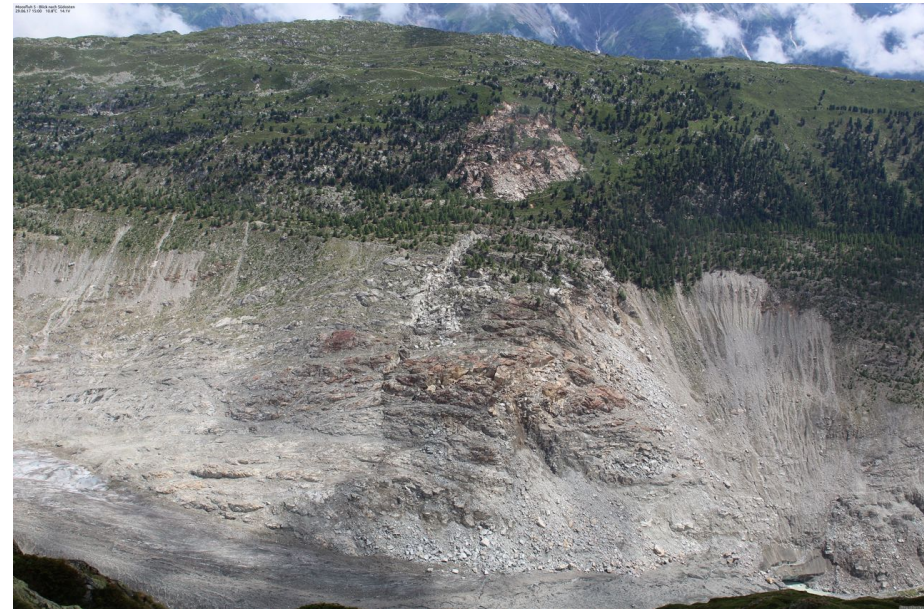
Thickness  $\approx 150$  m

Max velocity  
 $\approx 1$  m/day



# Aletsch

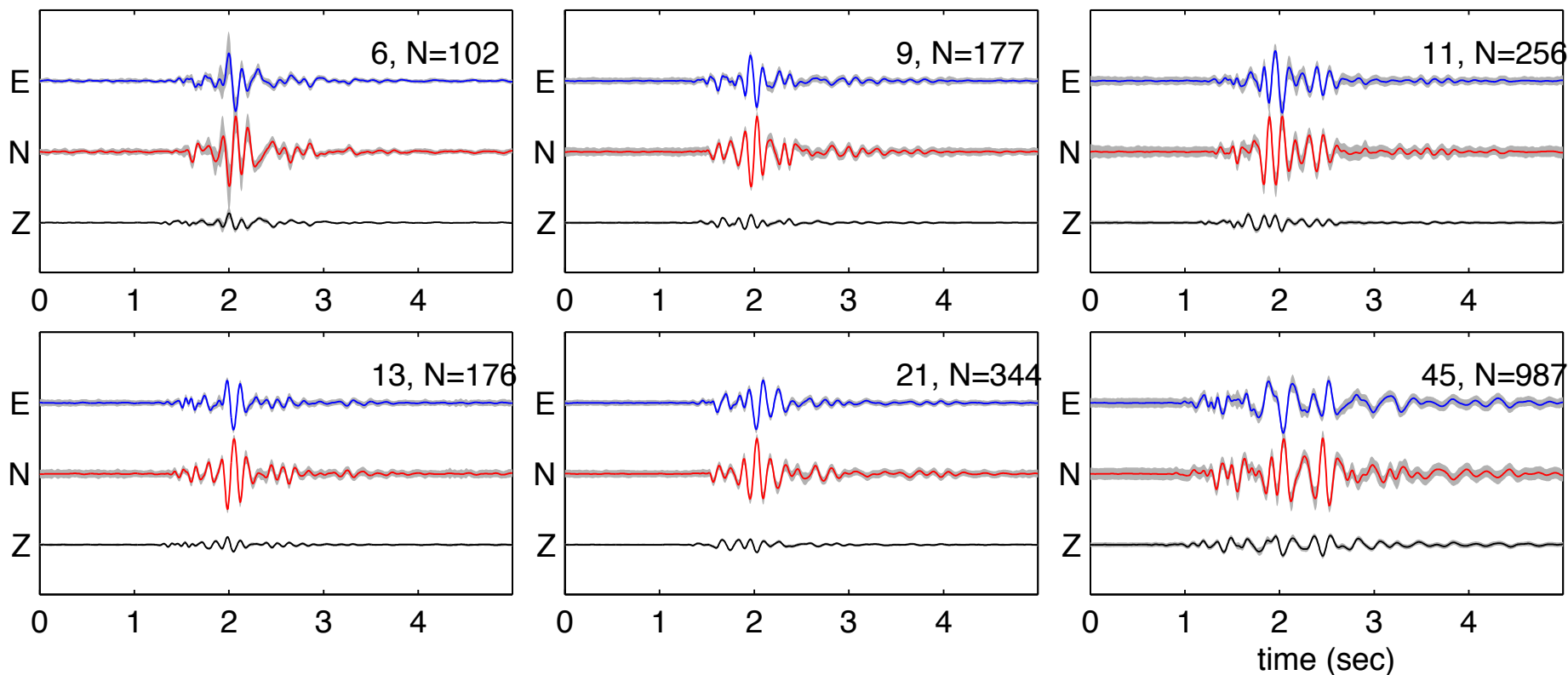
1 image / month  
march-september 2017





# Seismic signals

- > Duration  $\approx 2$  s, frequency  $\approx 5$  Hz
- > 30000 events detected by template-matching since 2016/10/3
- > 640 clusters with up to 1474 events



# Seismic activity and displacement rate

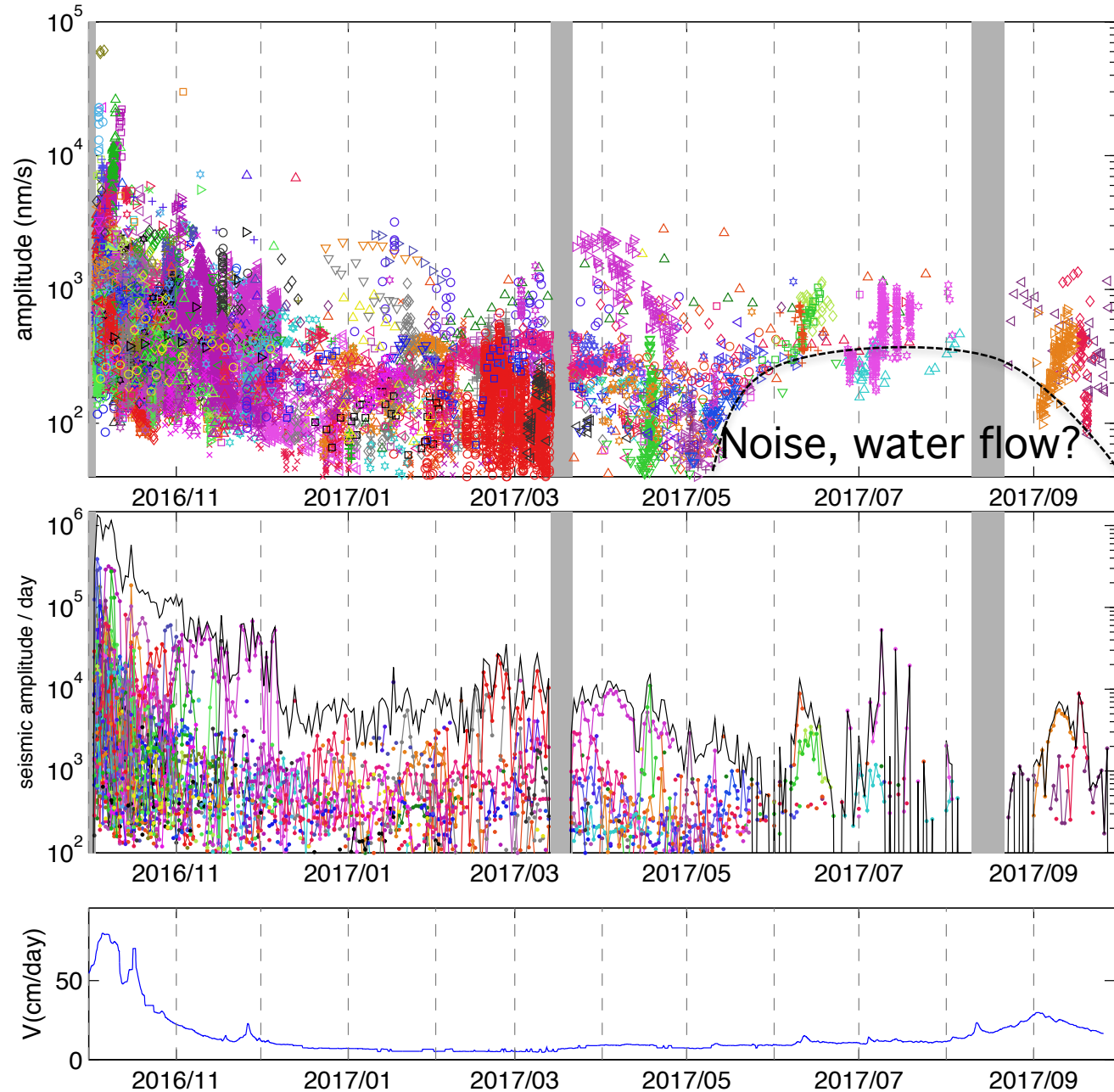
Selection of «repeaters»

Regular in time and size

# symbol for each  
cluster

Cumulated  
amplitude/day

Rockslide velocity  
(GPS data ETHZ)

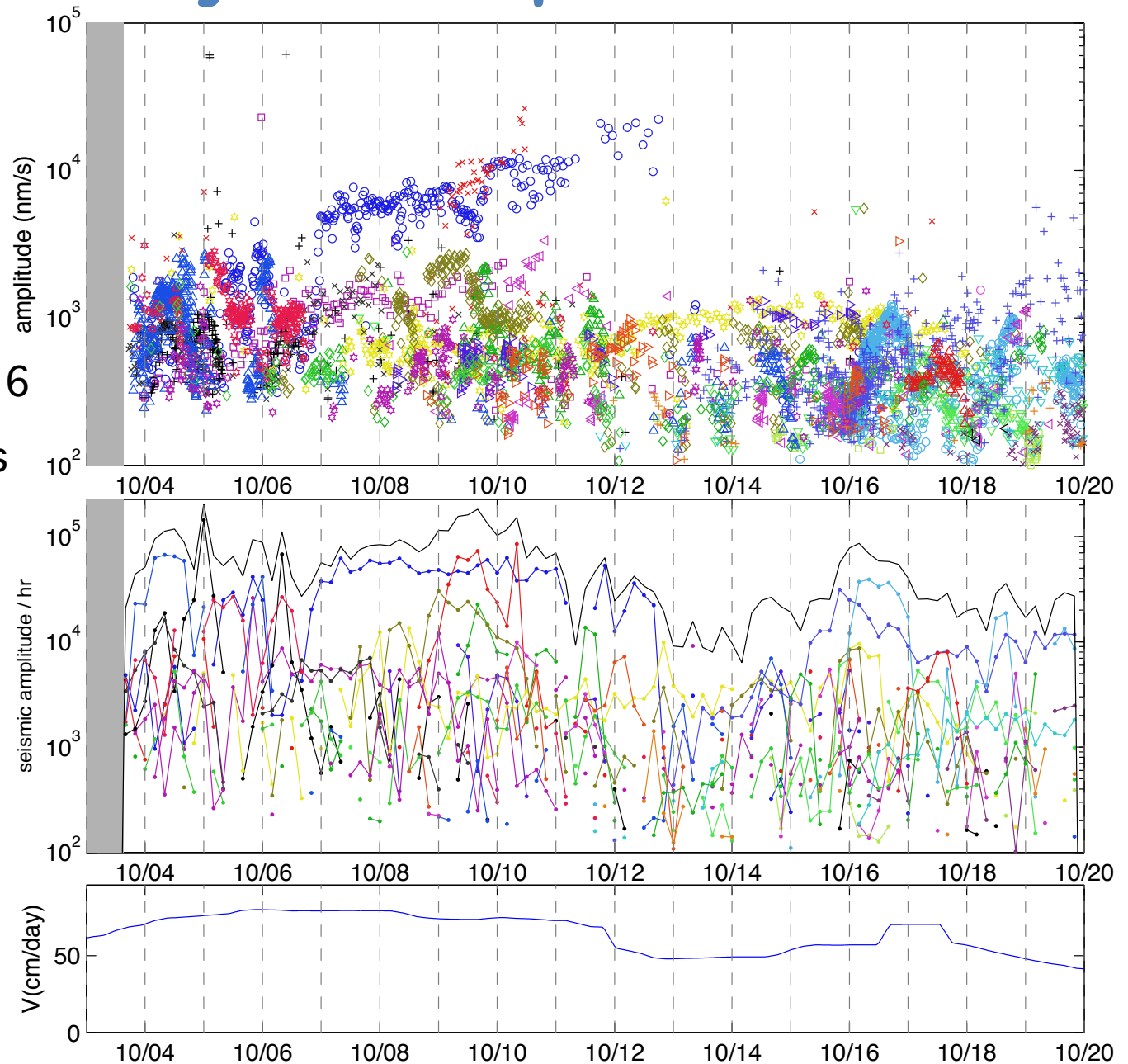


# Seismic activity and displacement rate

Zoom

3-20 October 2016

Largest sequences

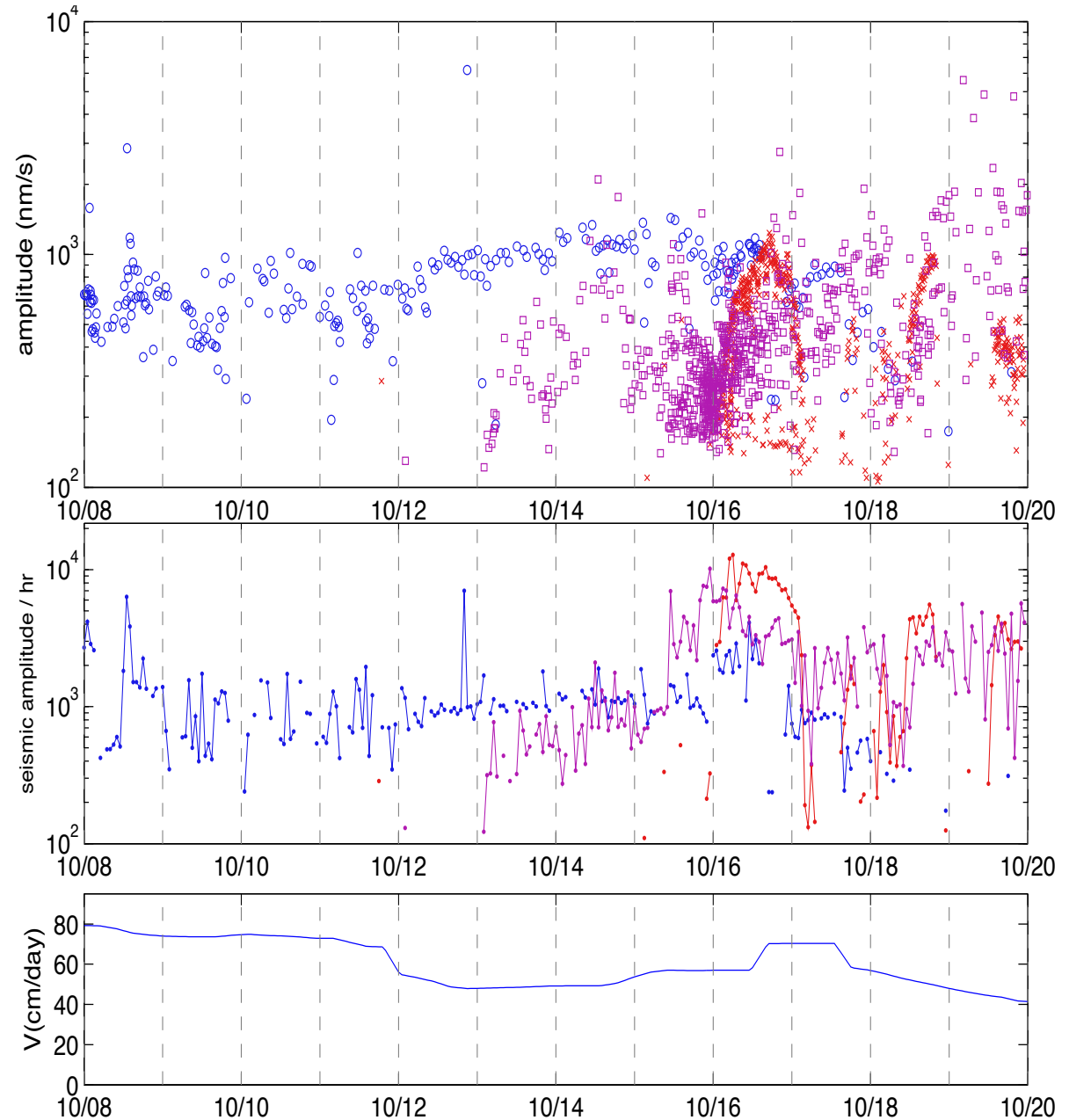


# Seismic activity and displacement rate

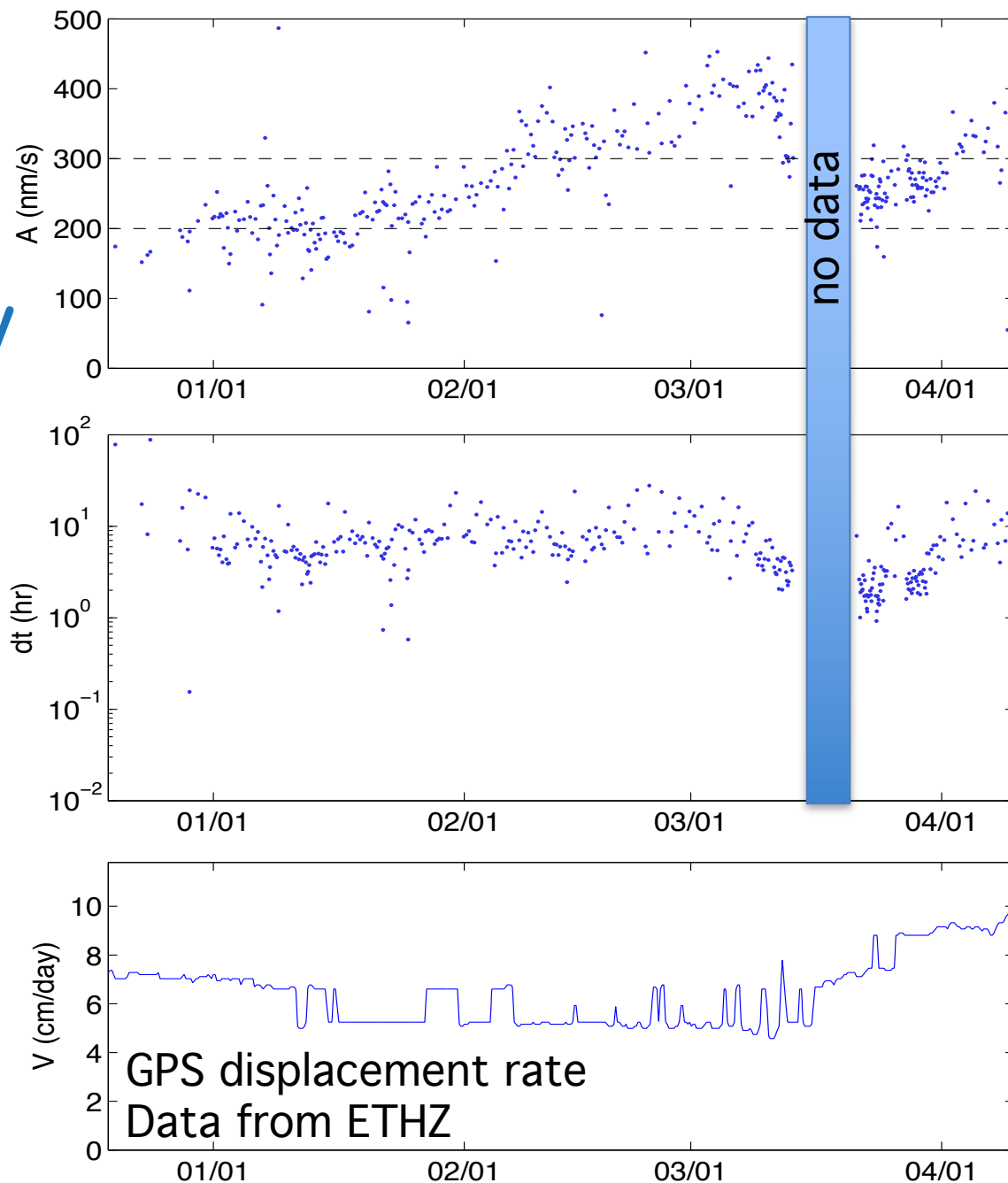
Zoom

10-20 October 2016

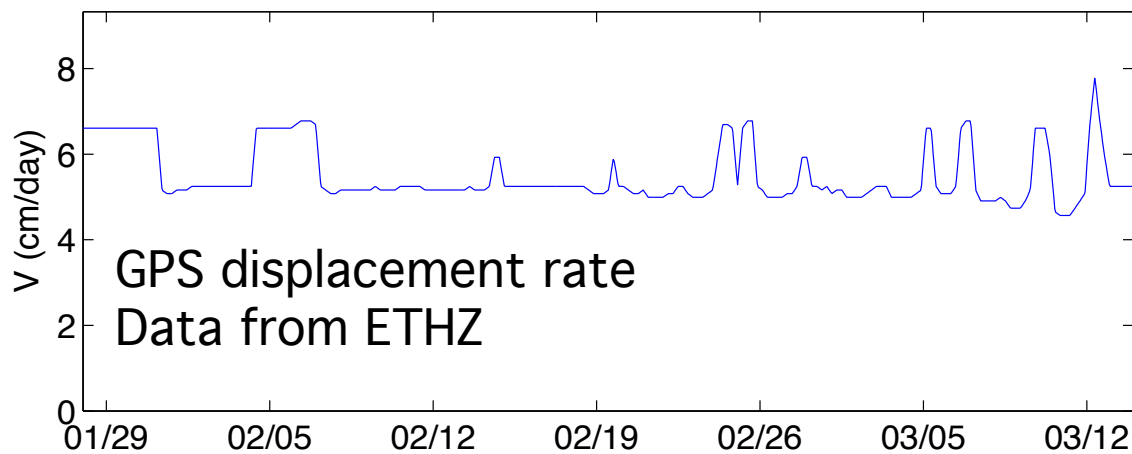
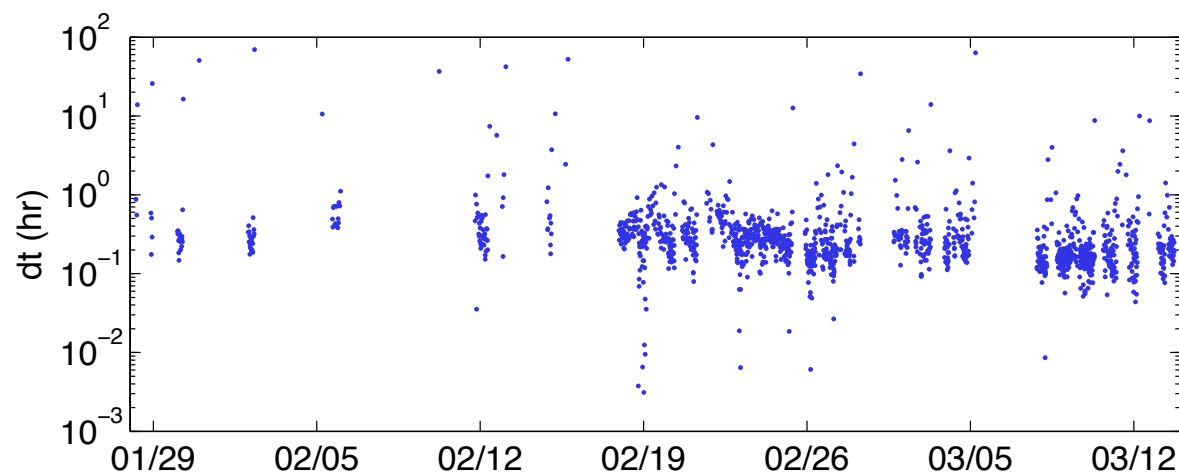
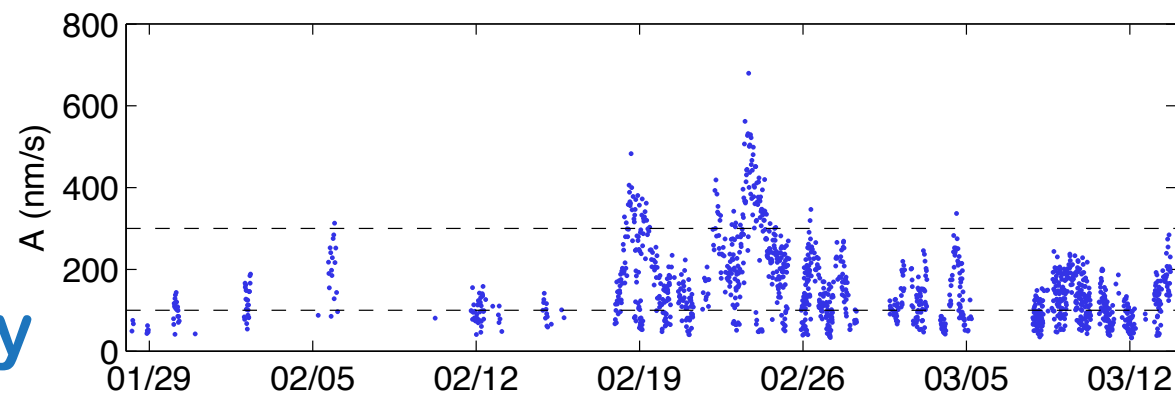
3 sequences



# Relation between peak amplitude, recurrence time and sliding velocity



# Relation between peak amplitude, recurrence time and sliding velocity



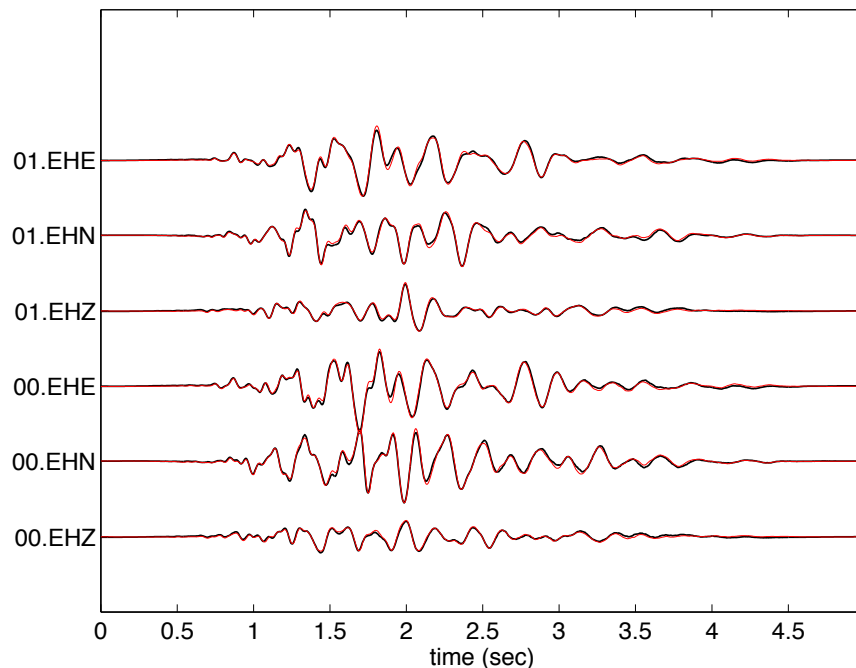
# Waveforms of small and large events

Black : stack of small events,  $A < 100$  nm/s and  $\langle A \rangle = 71$  nm/s

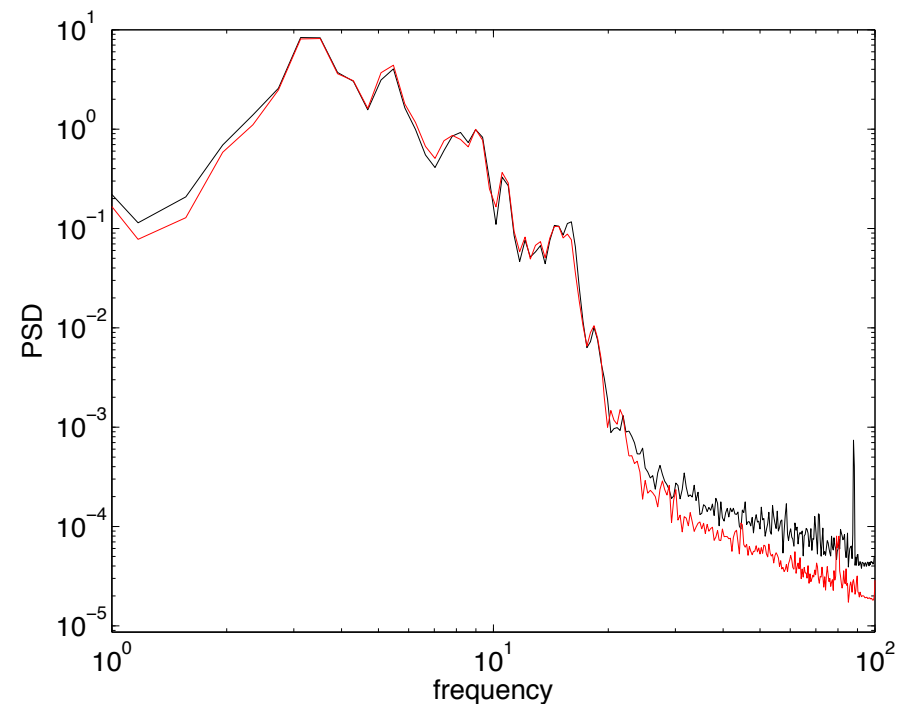
Red : stack of large events,  $A > 300$  nm/s and  $\langle A \rangle = 378$  nm/s

Magnitude ?  $<0$  ? Corner frequency  $> 1000$  Hz?

Stacked waveforms, normalized  
by peak amplitude  $A$ , no filter



Normalized spectrum  
Smaller events have relatively more  
energy for  $f > 20$  Hz

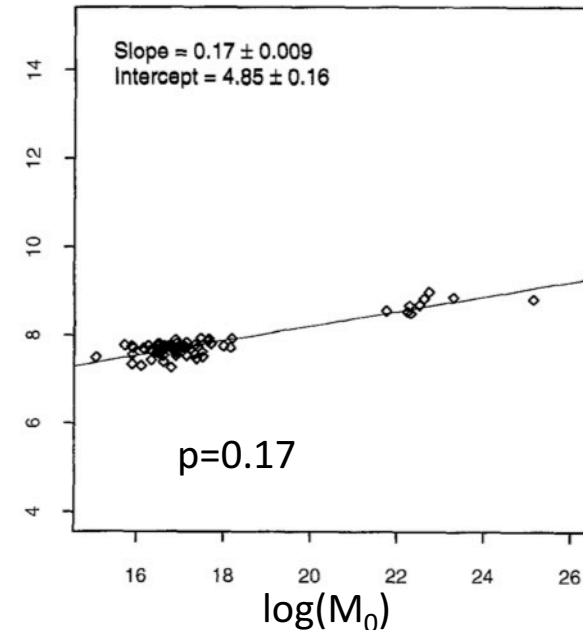
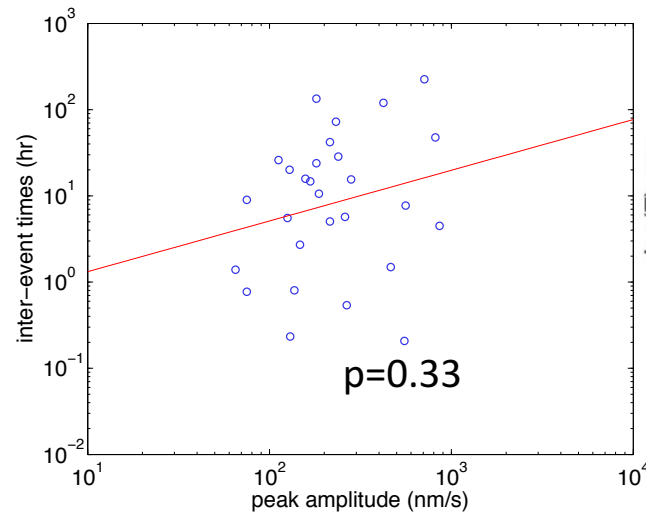
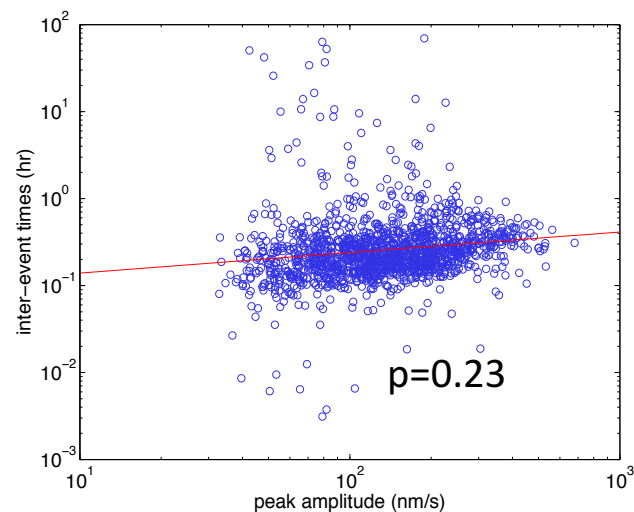


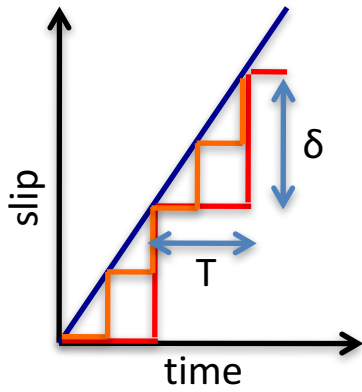


# Relation between size and recurrence time

Interevent times as a function of amplitude for :

- > 1 sequence of repeaters in Aletsch
- > Average values for the 28 most regular sequences with times between 2016/12/1 and 2017/9/1
- > Average for 62 sequences in Parkfield ( $p=0.17$ ; Nadeau and Johnston, 1998)





$$T \sim M_0$$

Constant  $L$  (Lengliné et al 2014)

No interseismic slip

Constant loading  $\Rightarrow T \sim \delta$

$$M_0 \sim \delta L^2 \sim \delta \sim T$$

$$T \sim M_0^{1/3}$$

Constant stress drop  $L \sim \delta$

No interseismic slip

$$M_0 \sim \delta L^2 \sim \delta^3 \sim T^3$$

$$T \sim M_0^{1/6}$$

> Nadeau and Johnson (1998)

no interseismic slip  $\Rightarrow T \sim \delta$

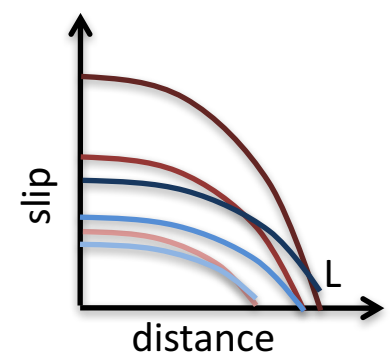
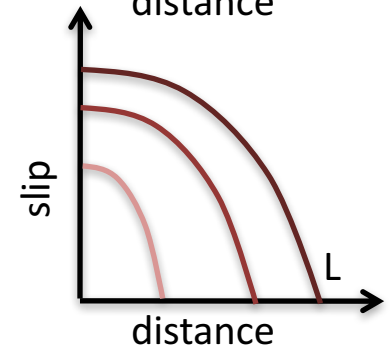
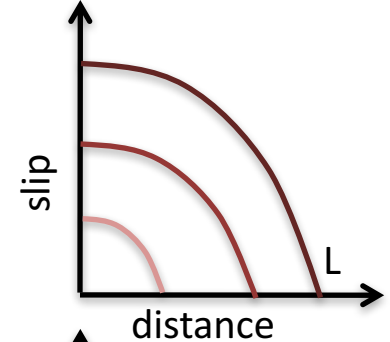
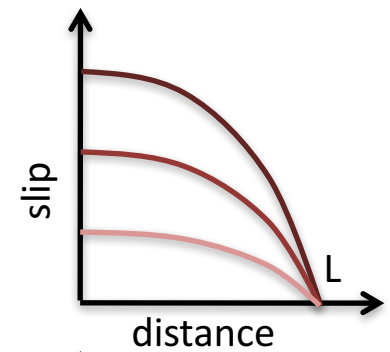
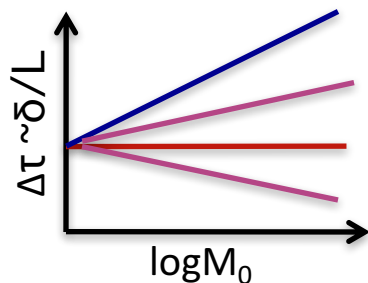
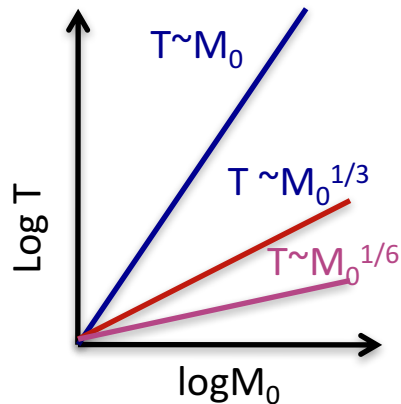
stress drop decreases with  $M_0$

$$M_0 \sim \delta^6 \sim T^6$$

> Cattania (2017)

interseismic slip  $\Rightarrow T \sim L$

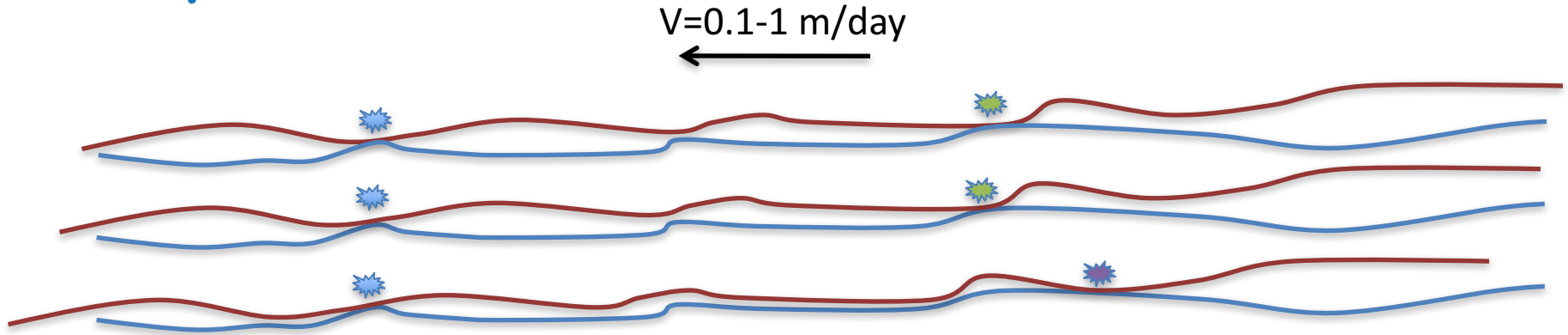
stress drop increases with  $M_0$



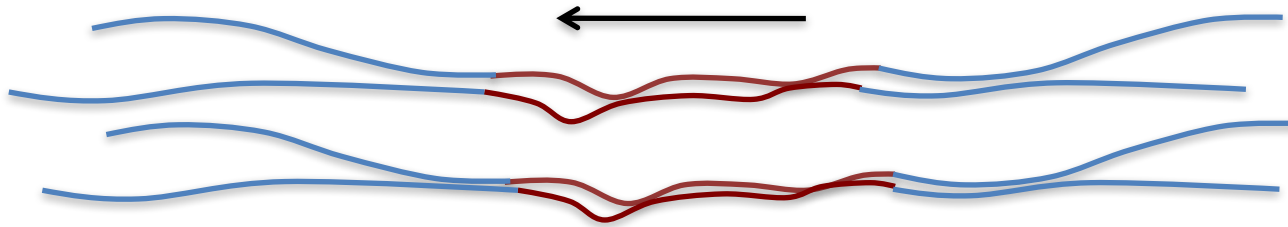
# Temporal variation of amplitude and recurrence times within each cluster

- > No correlation between # clusters
- > No correlation with precipitations, snow or temperature
- > No correlation with rockslide velocity
- > Variations dues to local properties of the interface?

# Asperities ...



At large scale, appearance and destructions of asperities with slip

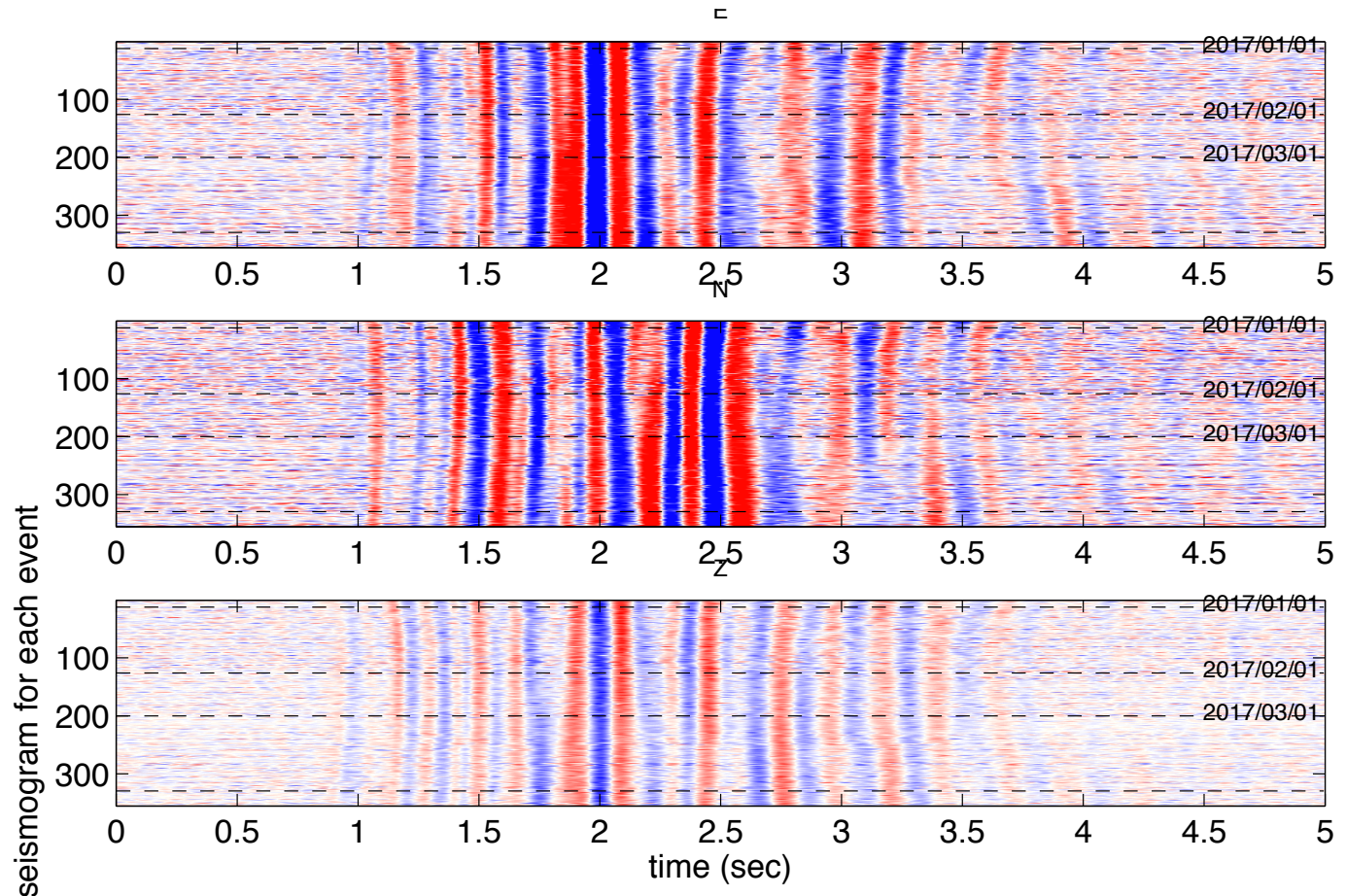


At the scale of one asperity ( $\approx 1 \text{ m}$ ),

a small slip can change the contact area and the stress on the interface  
yielding a change in amplitude and recurrence time?

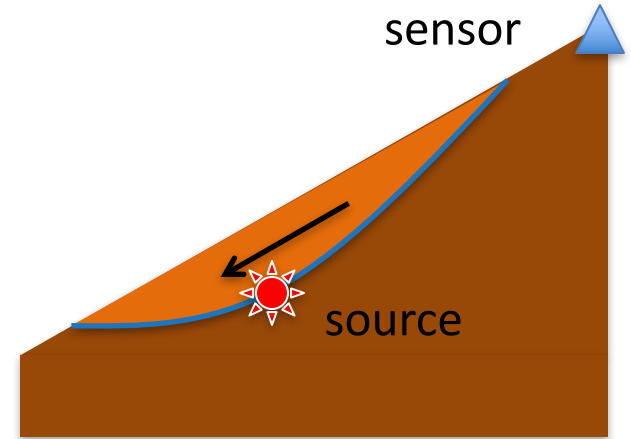
# Temporal changes in seismic waveforms

- > One sequence of repeaters with 355 events in 4 months
- > 1 line = 1 event, color ~ seismic ground velocity
- > 1 plot for each trace (Z,N,E)
- > Source moving or change in the medium?



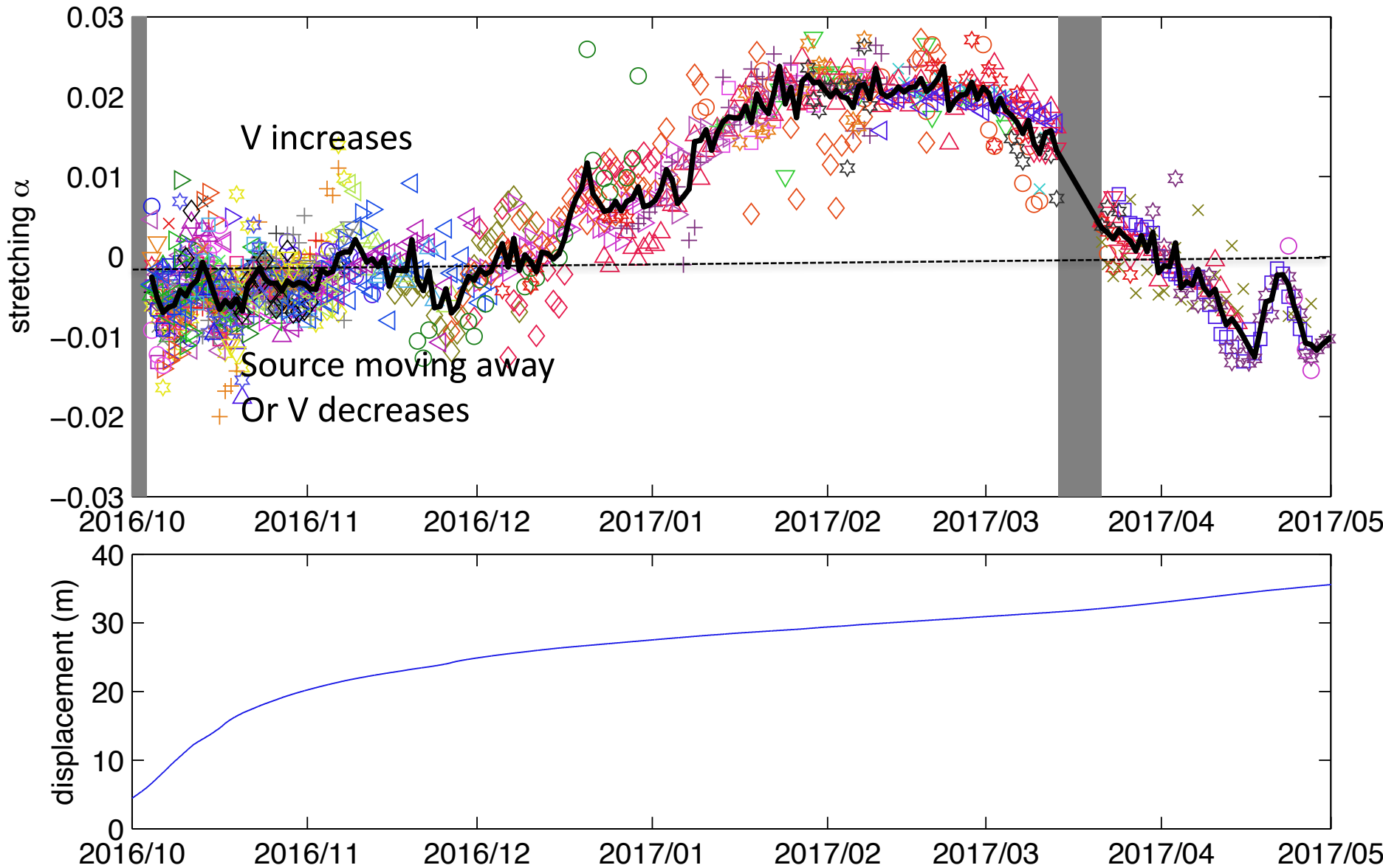
# Temporal variation of waveforms

- > Selection of « repeating » events
- > Shifting and stretching of signals
- > Average over 1 day => waveform  $y_i(t)$
- > Correlation with the stack of all events  $y_t$
- > Find  $dt_i$  and  $\alpha_i$  that maximize the correlation between  $y_i((t+dt_i)*(1+\alpha_i))$  and  $y_t$
- > Fixed source =>  $\alpha = dV/V$  : variation of seismic wave velocities
- > Source moving with rockslide =>  $\alpha \sim \text{displacement}$  and  $\alpha < 0$
- > Align all curves for # clusters by minimizing residuals and average



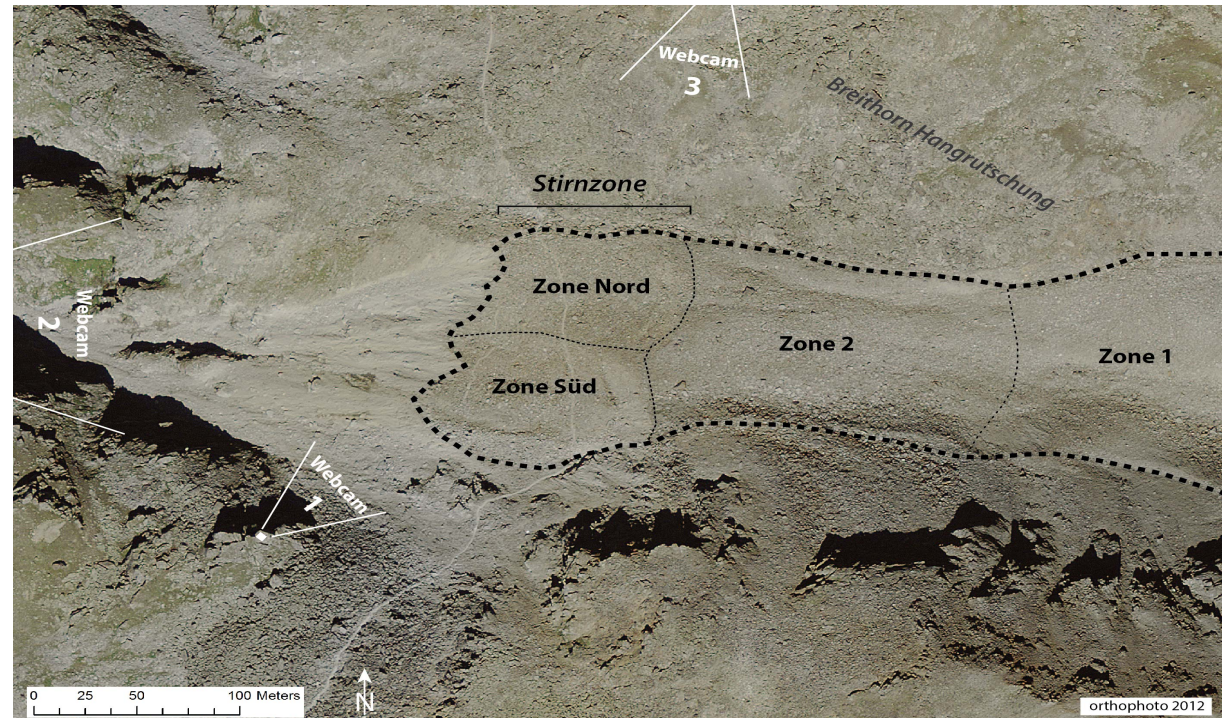
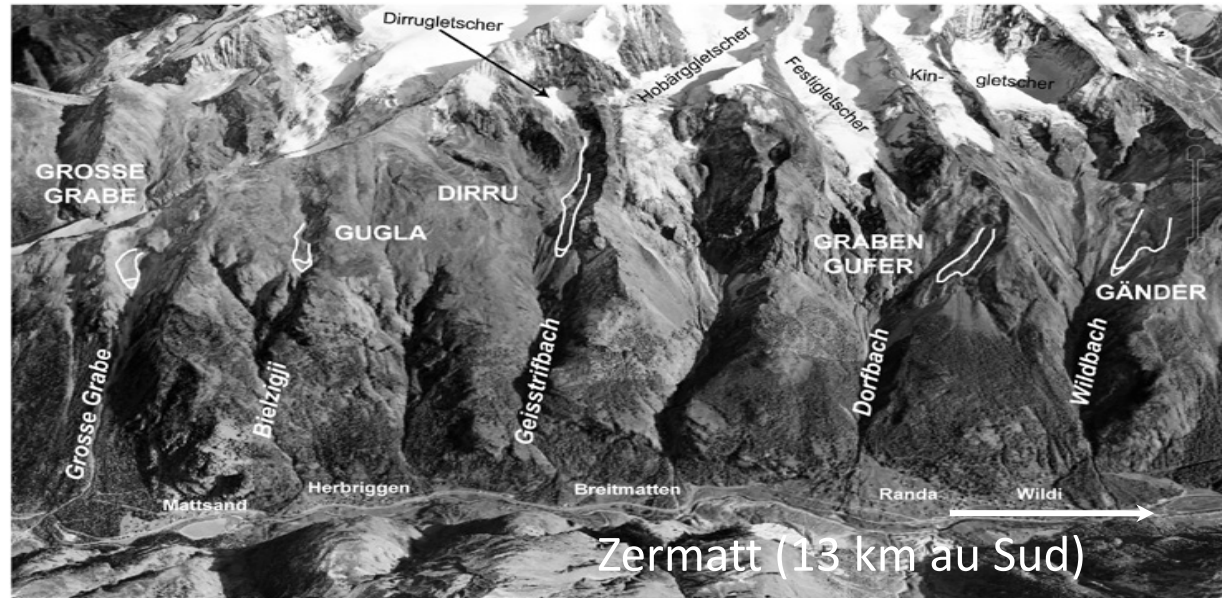


# Temporal variation of waveforms



# Gugla

Rock-glacier  
in Switzerland  
(Valais near Zermatt)  
Elevation 2600-3000 m  
length ~600 m  
Width 130 m





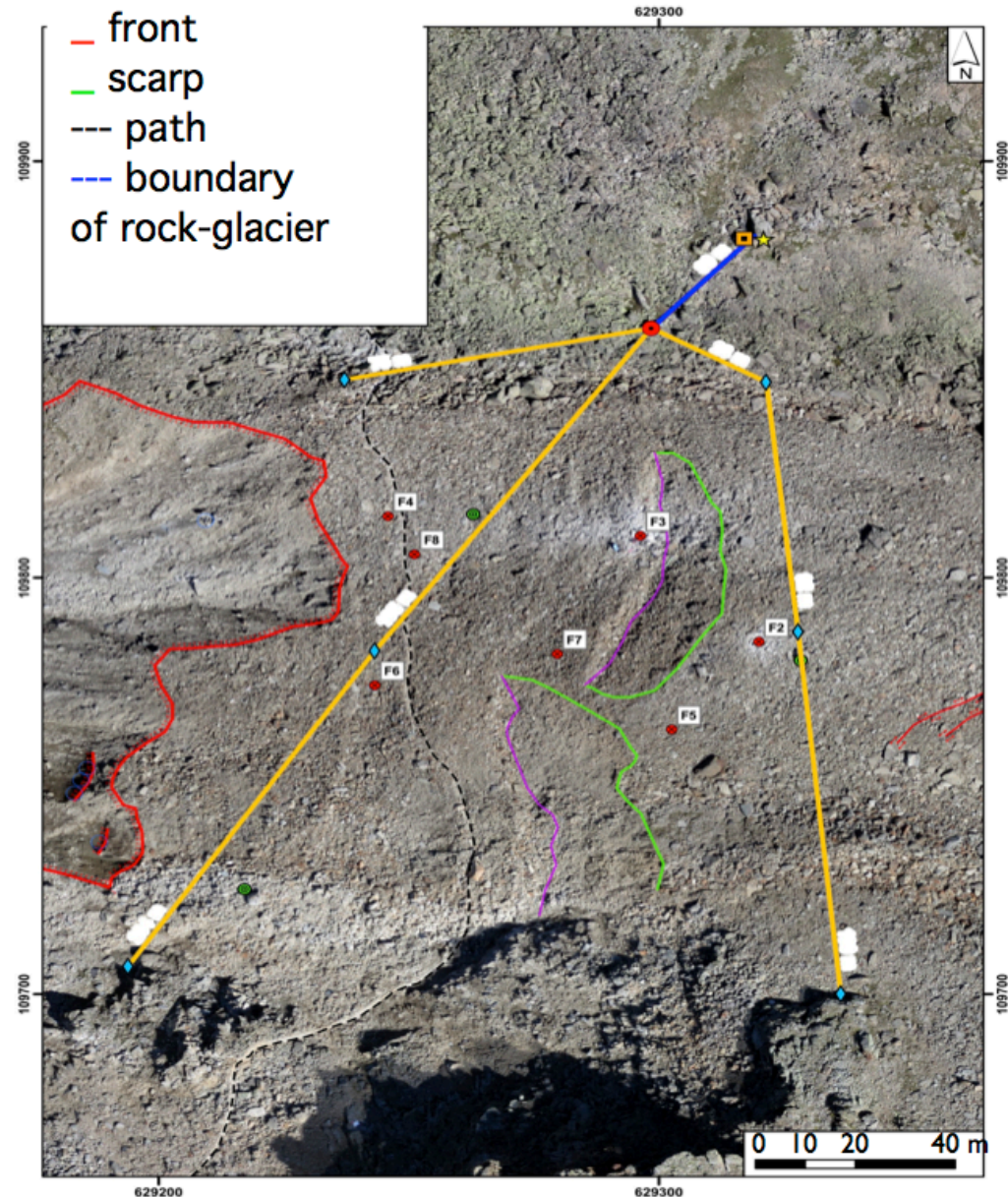
# Gugla - Instrumentation

## Seismology (ISTerre)

- ◇ 6 1C 2Hz seismometers
  - ▣ digitizer (continuous 200 Hz)
- installed Oct 2015

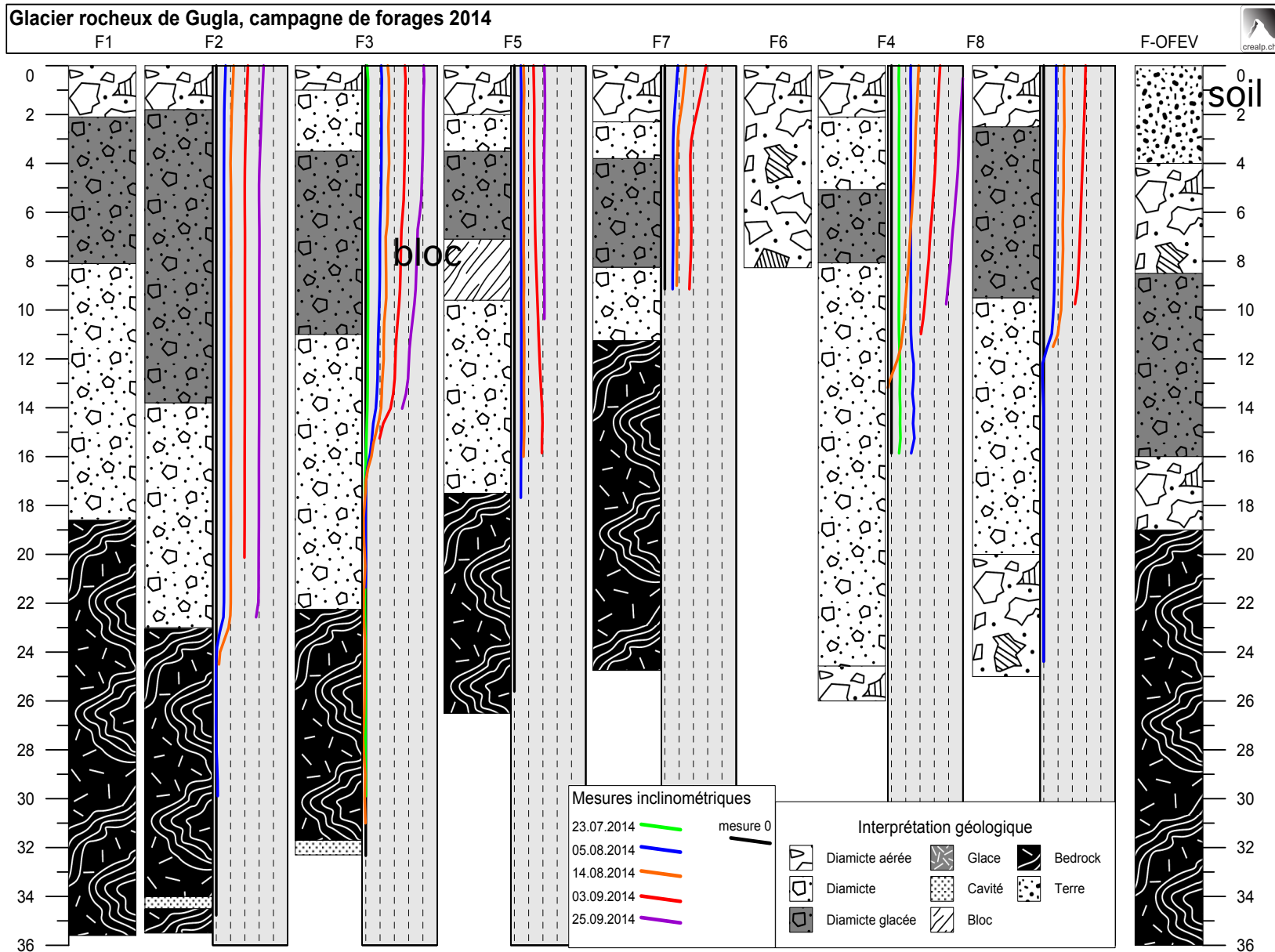
## Displacement (Univ. Fribourg and ETHZ):

- ⊙ GPS : 60 points measured (twice) every yr1 continuous GPS since 2007
  - ⊗ 7 boreholes drilled in 2014 by CREALP with inclinometers and temperature
- 2 webcams



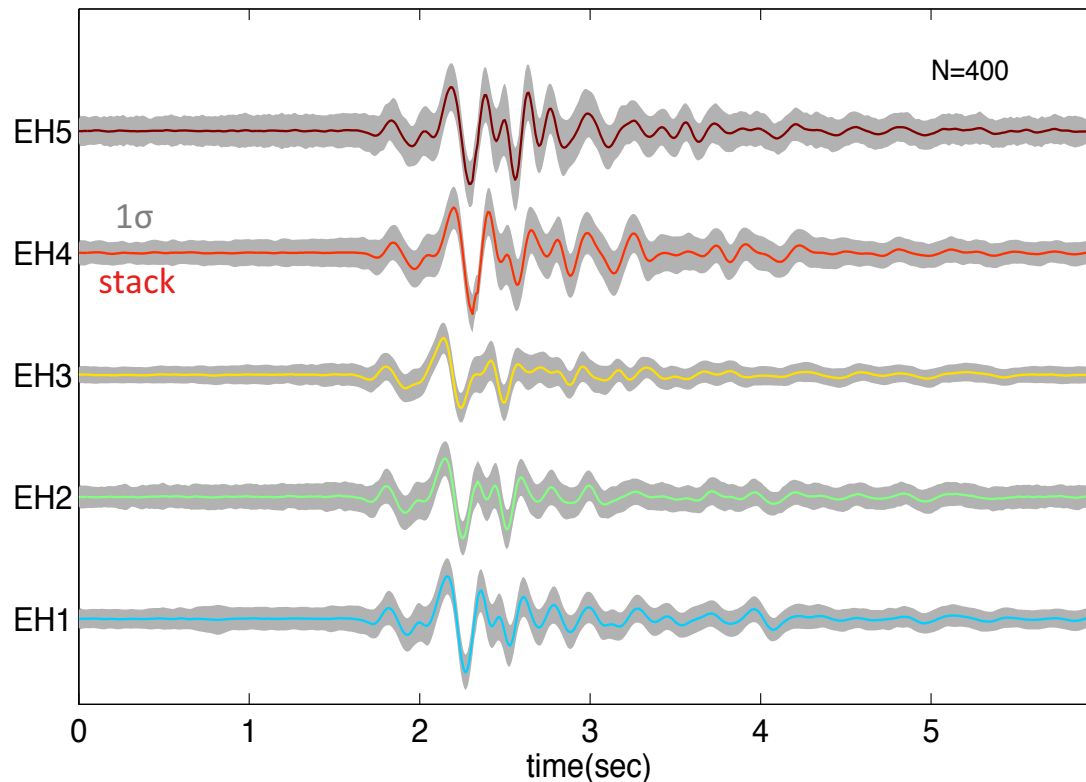
# Gugla – Boreholes and inclinometric data

Data from CREALP and OFEV



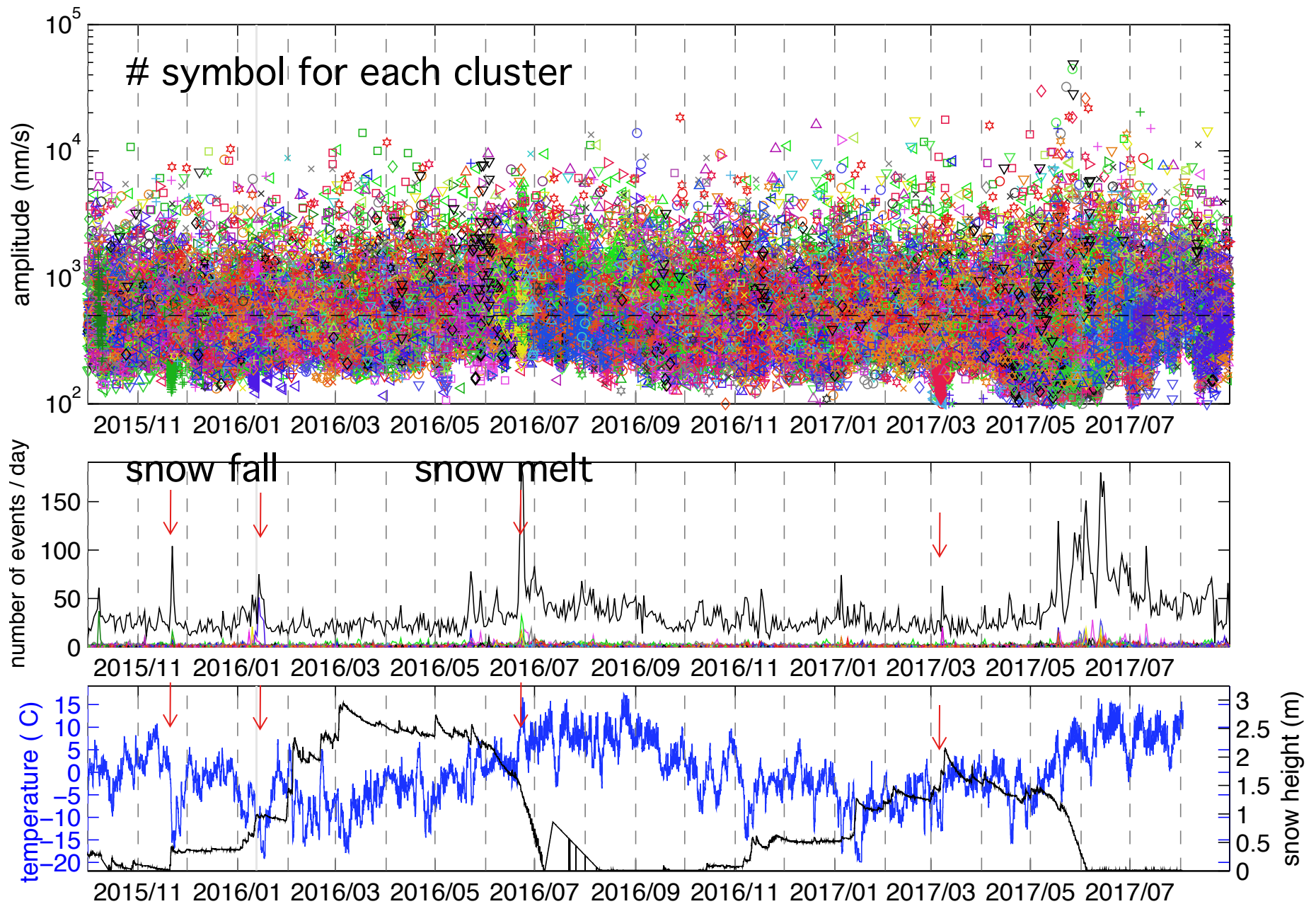
# Detection and classification of events

- > selection of short ( $\sim 5$ s ) and low frequency events ( $\sim 5$ Hz)
- > detection by template-matching
- >  $\approx 257$  clusters with up to 1100 events since 2015/10/1
- > good correlation between sensors and between events





# Seismic activity





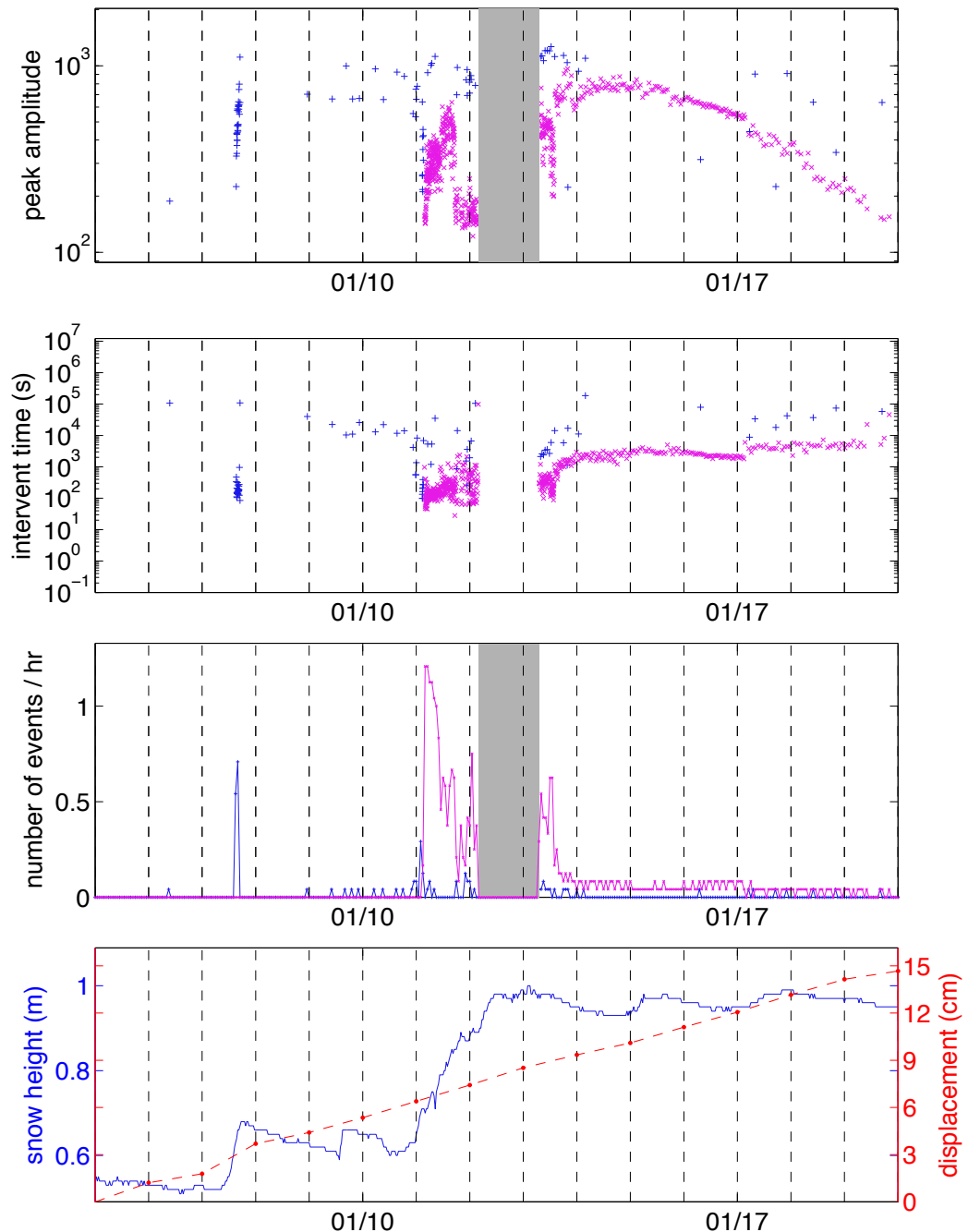
# Repeaters triggered by snow falls

Example for January 2016

Progressive change of amplitude and recurrence times

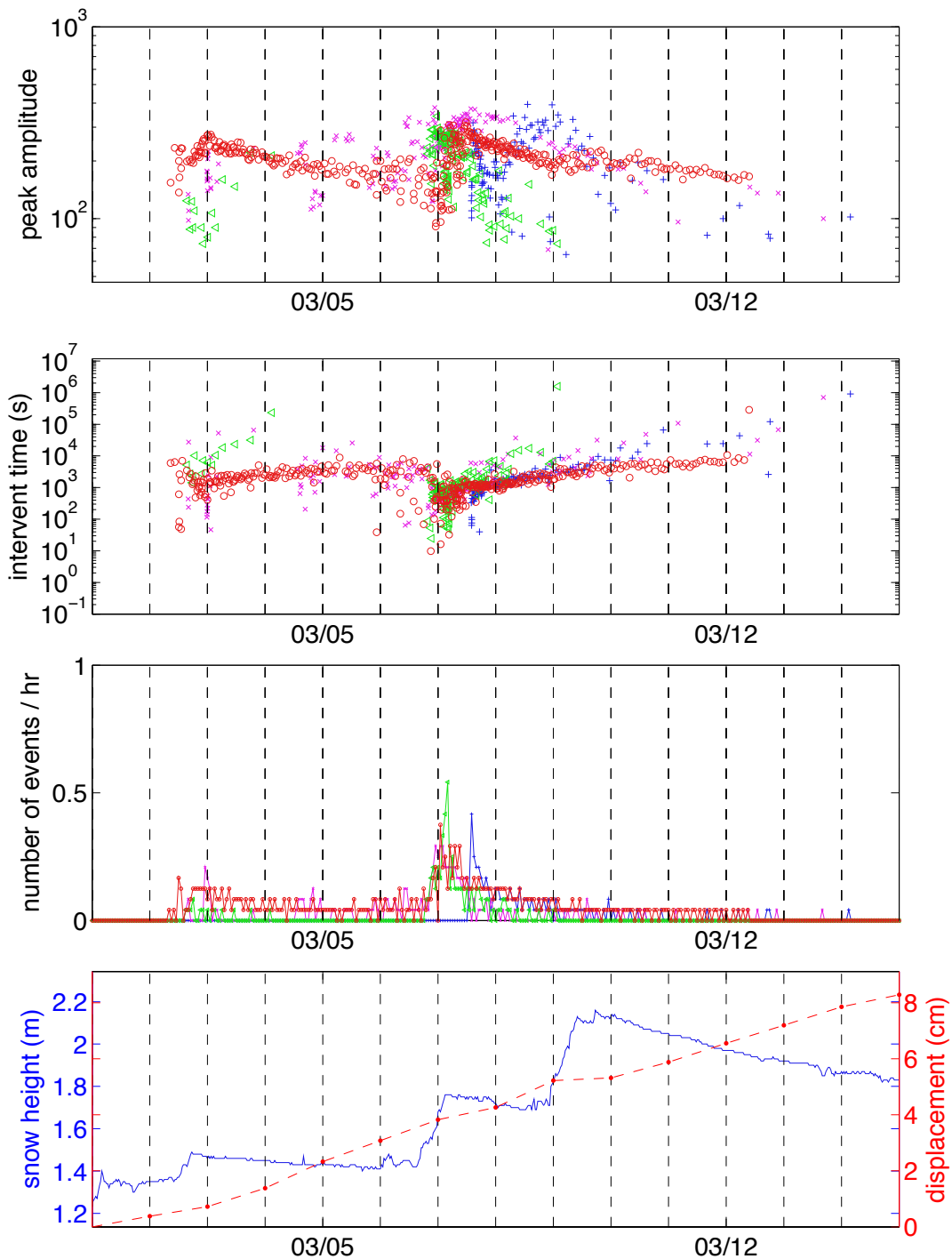
2 clusters activated

Almost no delay between snowfall and seismicity



# Repeaters triggered by snow falls

Example for March 2017

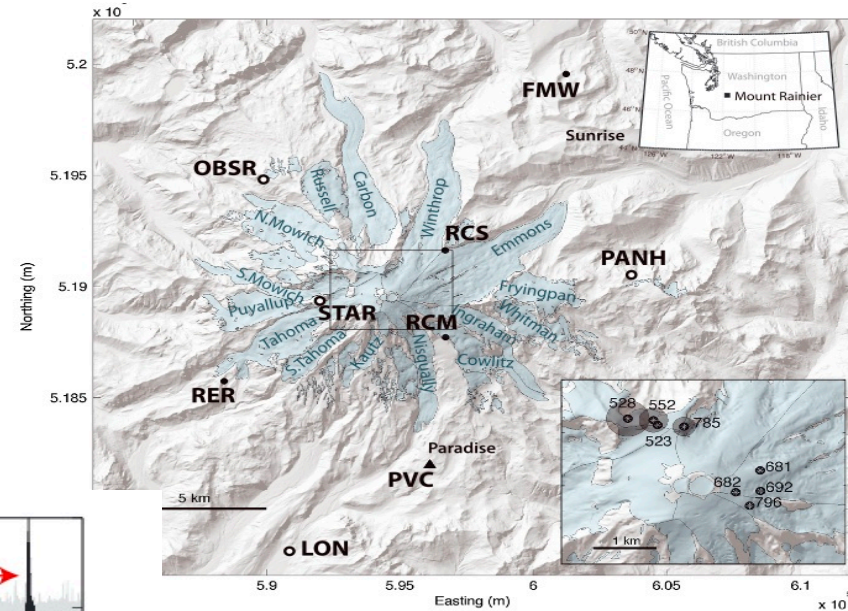
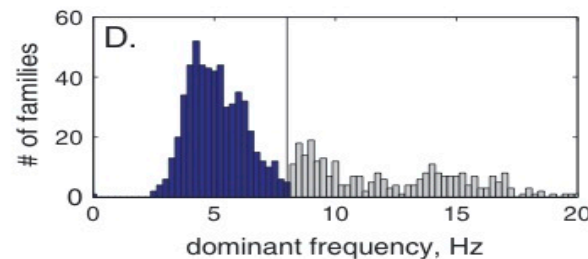
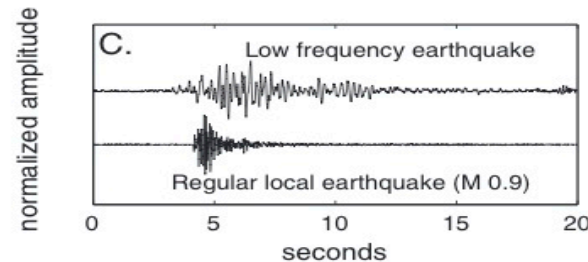
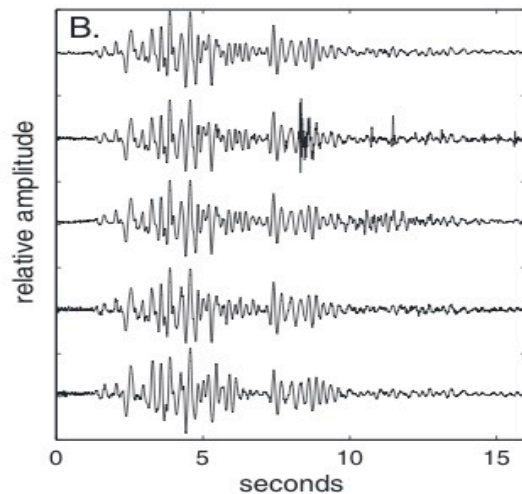
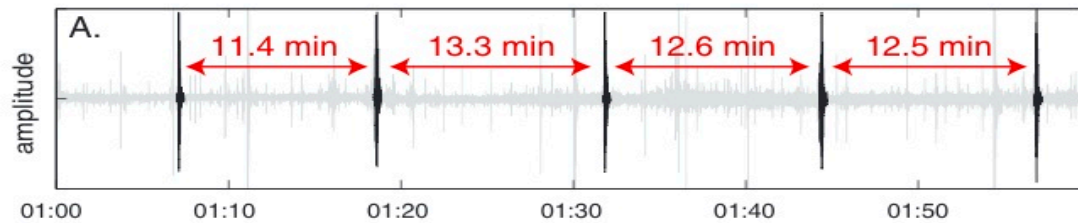


# Comparaison : Mount Rainier

Glacier-covered volcano

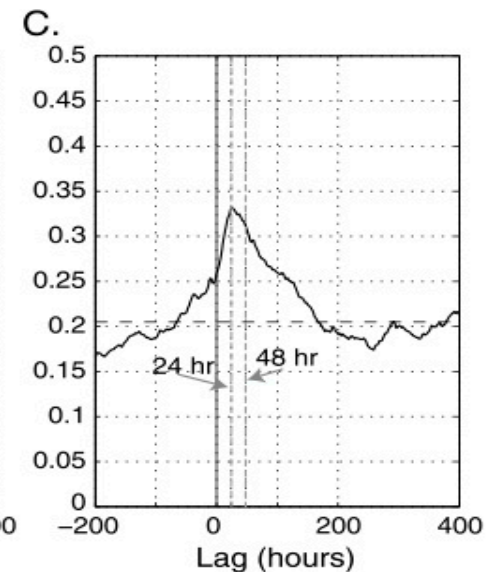
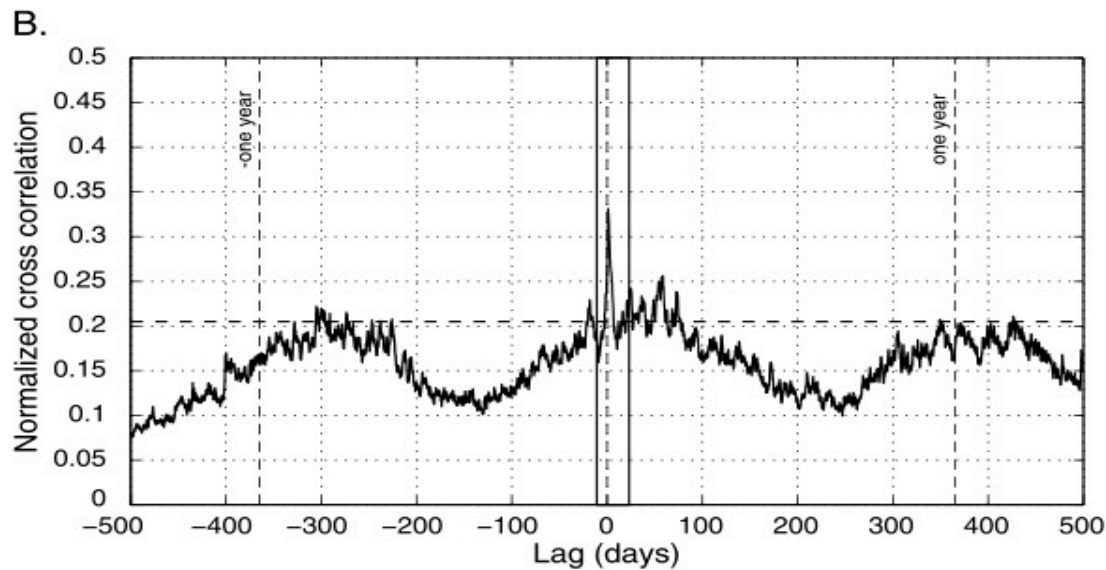
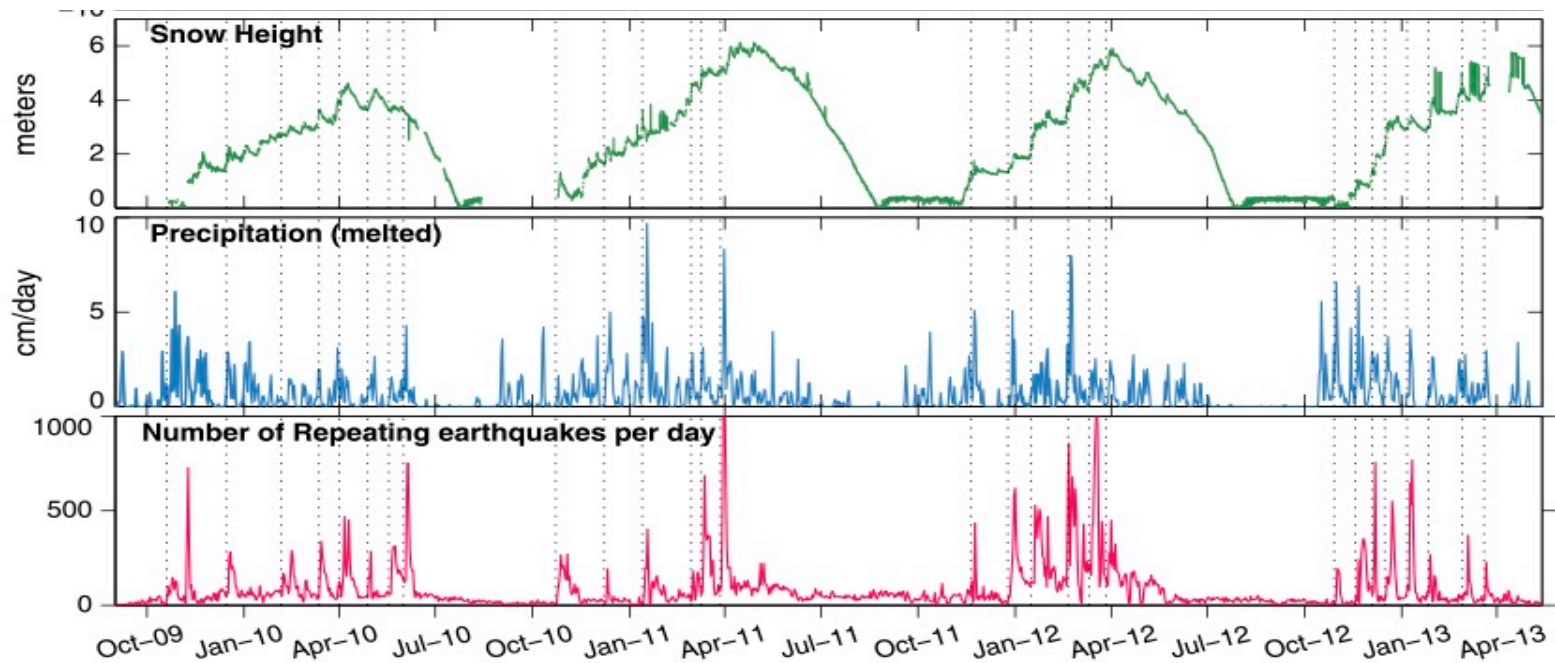
Multiplets of low frequency events( $\approx 5$  Hz)  
repeating every  $\approx 10$  mn

(Thelen et al 2013, Alstadt et al 2014)



# Comparison : Mount Rainier (US)

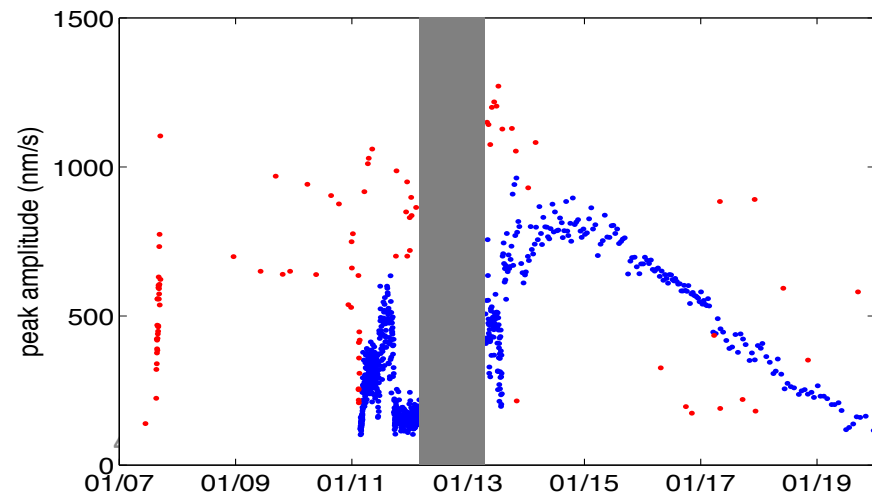
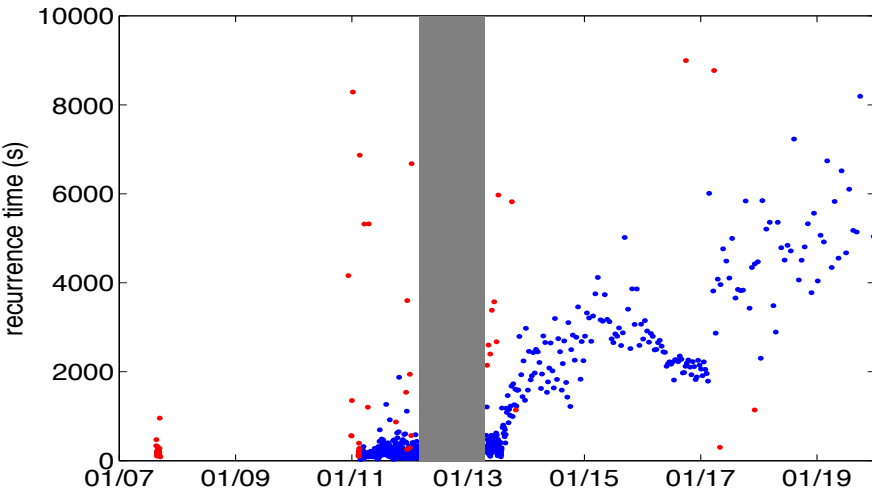
Alstadt et al 2014



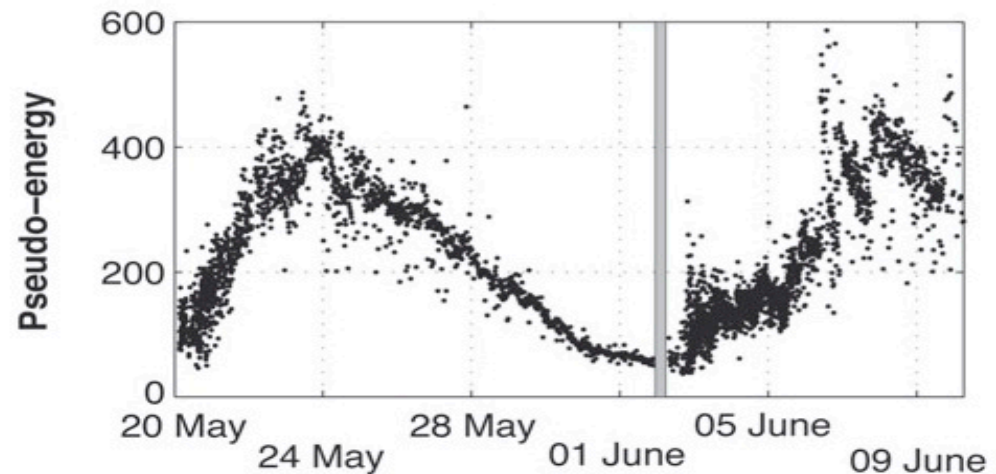
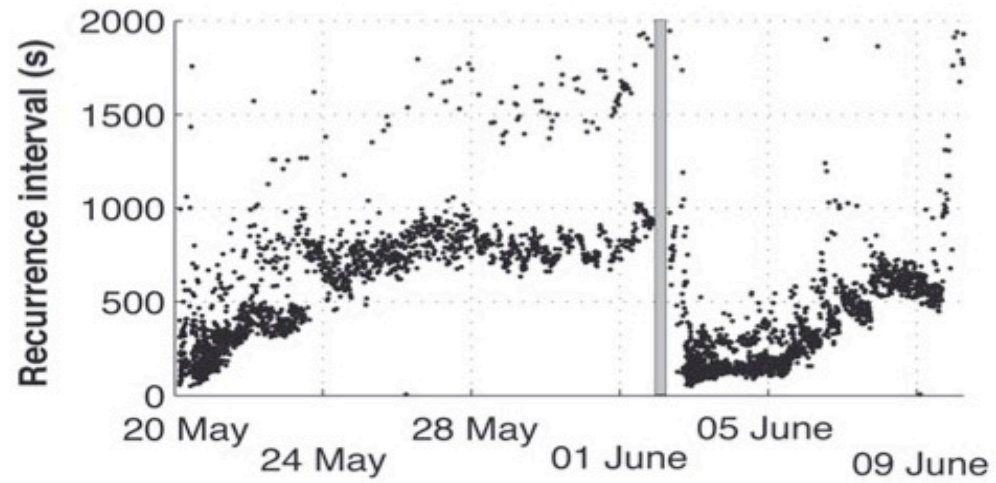
# Gugla / Mount Rainier

Progressive variations of amplitude and recurrence time

Gugla, January 2016



Mount Rainier (Allstadt et al 2014)





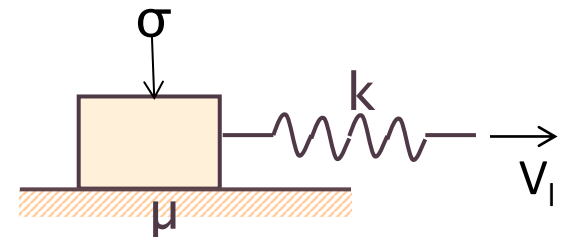
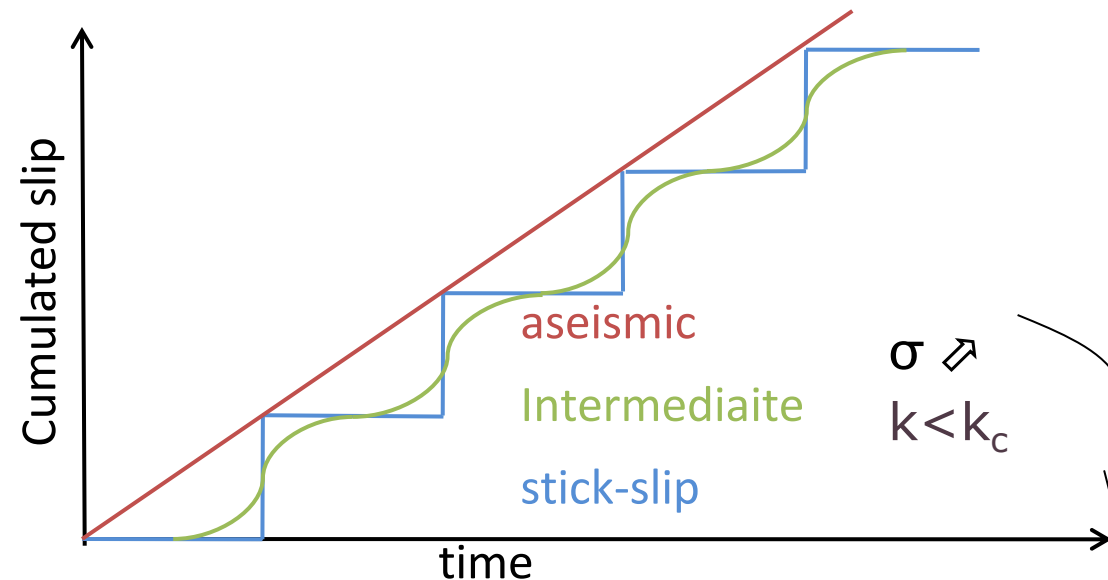
# Triggering mechanism

Allstadt et al (2014) : increase of normal stress due to snow weight  $\sim 0.1\%$   
+ redistribution, pressurisation of water in sub-glacial cavities  $\Rightarrow$  lubrication  
and slip acceleration

In summer : drainage network more developed  $\Rightarrow$  no over-pressure

In Gugla : very cold  $\Rightarrow$  no melt and no acceleration

Transition aseismic  $\Rightarrow$  seismic induced by increased normal stress?



# Conclusions (1/2)

≈ regular sequences of repeating icequakes and landquakes

But most events are random or clustered

Progressive variations of amplitude and recurrence times

No (systematic) correlation with sliding rate

No correlation and no interaction between # sequences

Weak correlation between recurrence time and size

## Conclusions (2/2)

- ⇒ Viscous deformation of ice or creep within the asperity
- ⇒ Variability mostly due to local changes of the interface
- ⇒ Landquakes are not (always) precursors of failure

Progressive changes in waveforms, mostly influenced by changes in the medium

Triggering by snowfall (Gugla, Rainier) : transition aseismic => seismic slip due to increased in normal stress?

# Perspectives

- > Observations :
  - add more seismometers in Alestch and Gugla in order to better locate and characterize repeaters
- > Numerical modeling and laboratory experiments
  - Triggering by an increase in normal stress
  - Temporal variations in seismic moment and recurrence rate
  - Transition seismic  $\Leftrightarrow$  aseismic
  - Transition random  $\Leftrightarrow$  regular
  - Scaling laws of slip, rupture length and recurrence times

# Aletsch instability

[M Sartori, October 2016]

